

# Solar irradiance variability and climate

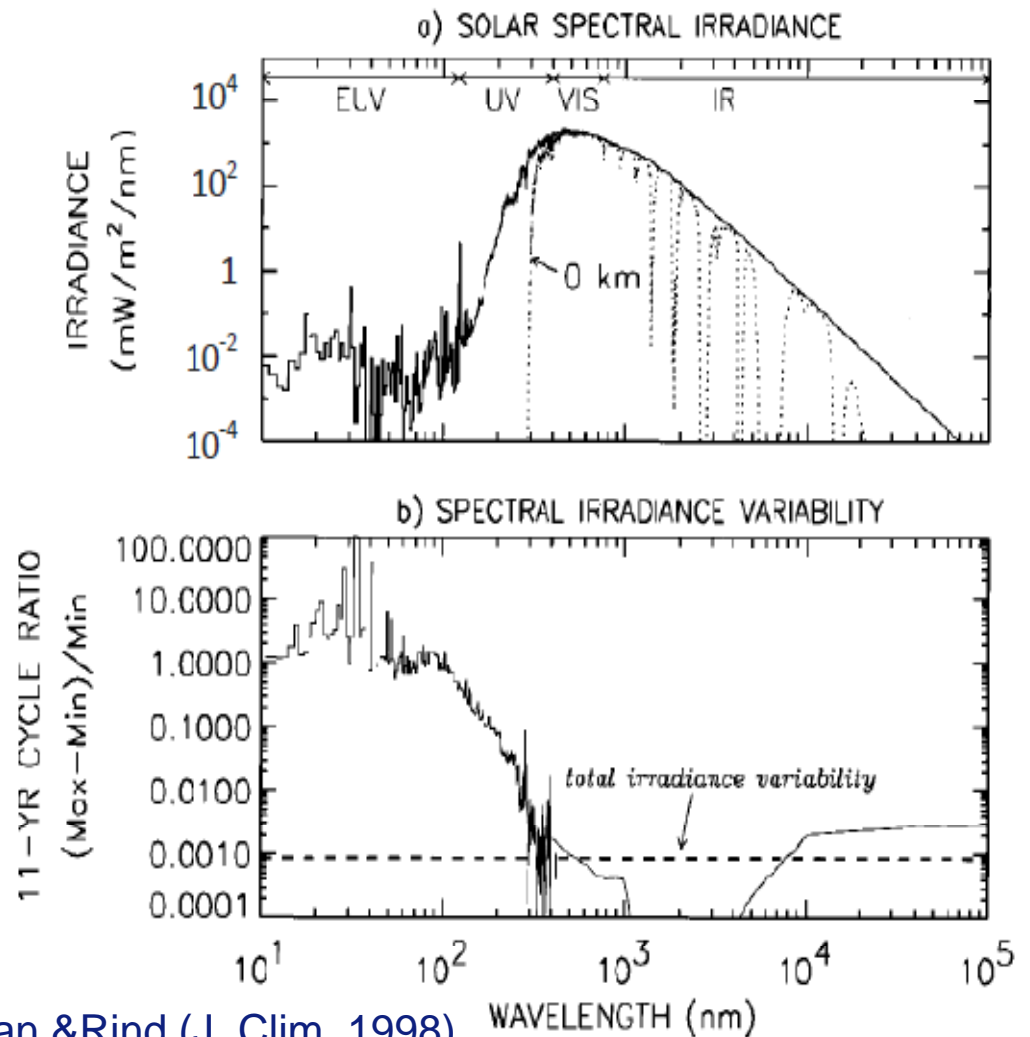
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Imperial College London, UK

Eddy Symposium Aspen CO 22-24 Oct 2010

# Background

- Evidence for solar influence on climate difficult to explain by small variations in total solar radiative output.
- UV variations (fractionally) much larger.



Lean & Rind (J. Clim., 1998)

- Visible wavelength radiation mainly reaches Earth's surface.
- UV radiation mainly absorbed in stratosphere.

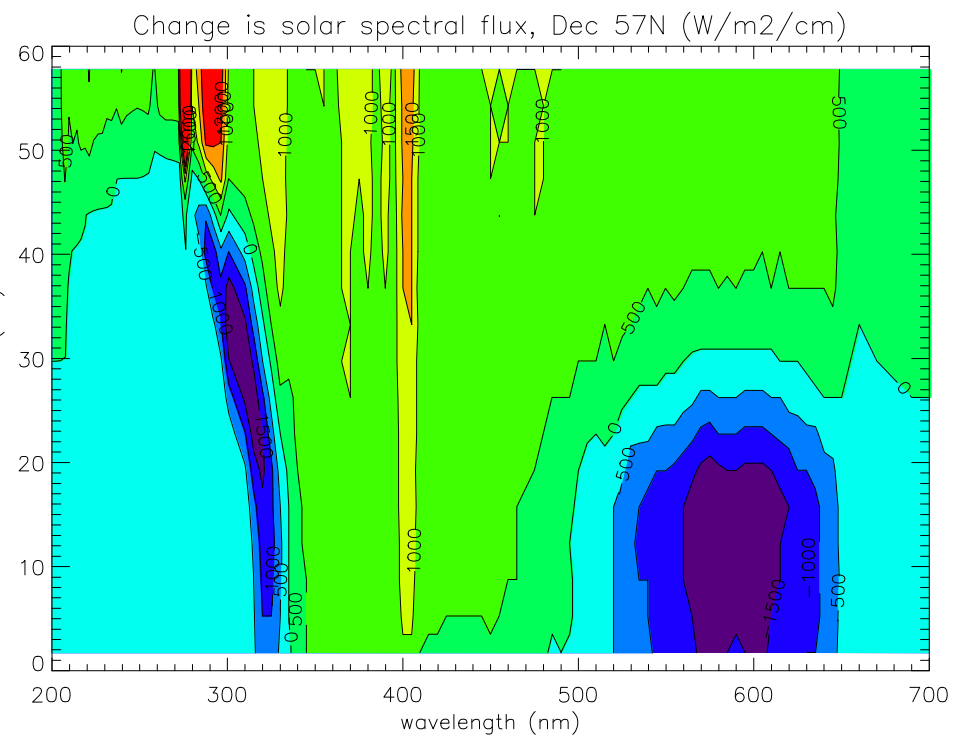
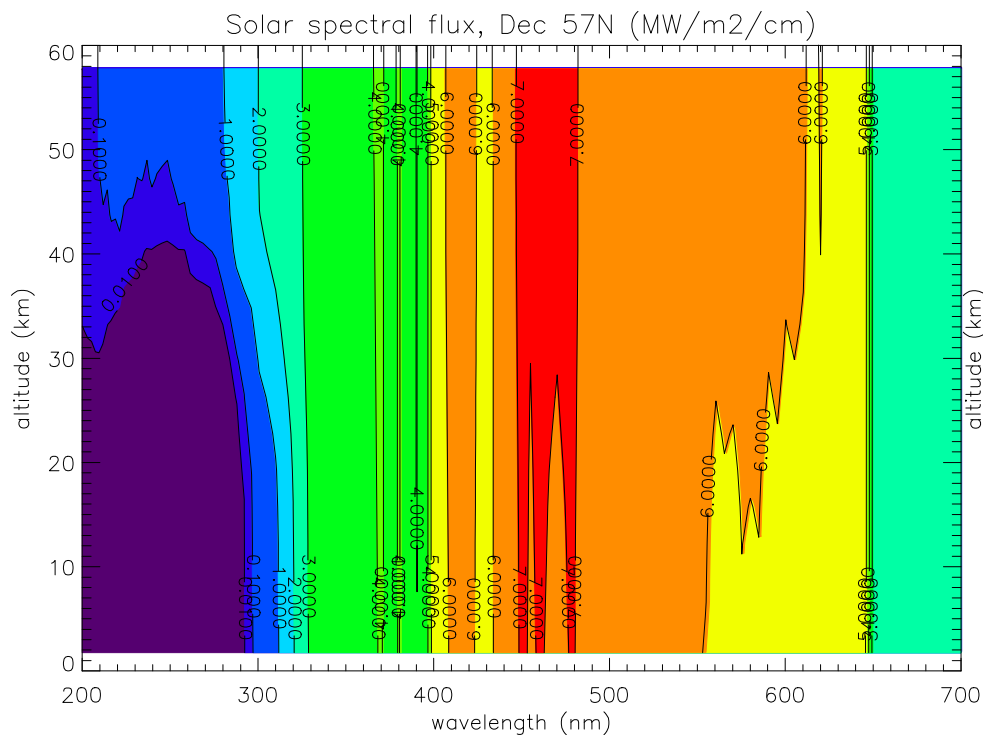
# Background:

## Solar cycle variation in spectral irradiance

solar min

(Dec, 57N)

max-min



Haigh (Nature, 1994)

Solar irradiance reaching lower atmosphere depends on zenith angle and on response in stratospheric ozone.

# Solar forcing of climate: radiative mechanisms

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“Top-down” via UV heating the stratosphere

and/or

“Bottom-up” via (visible) radiation warming surface ?

# Outline

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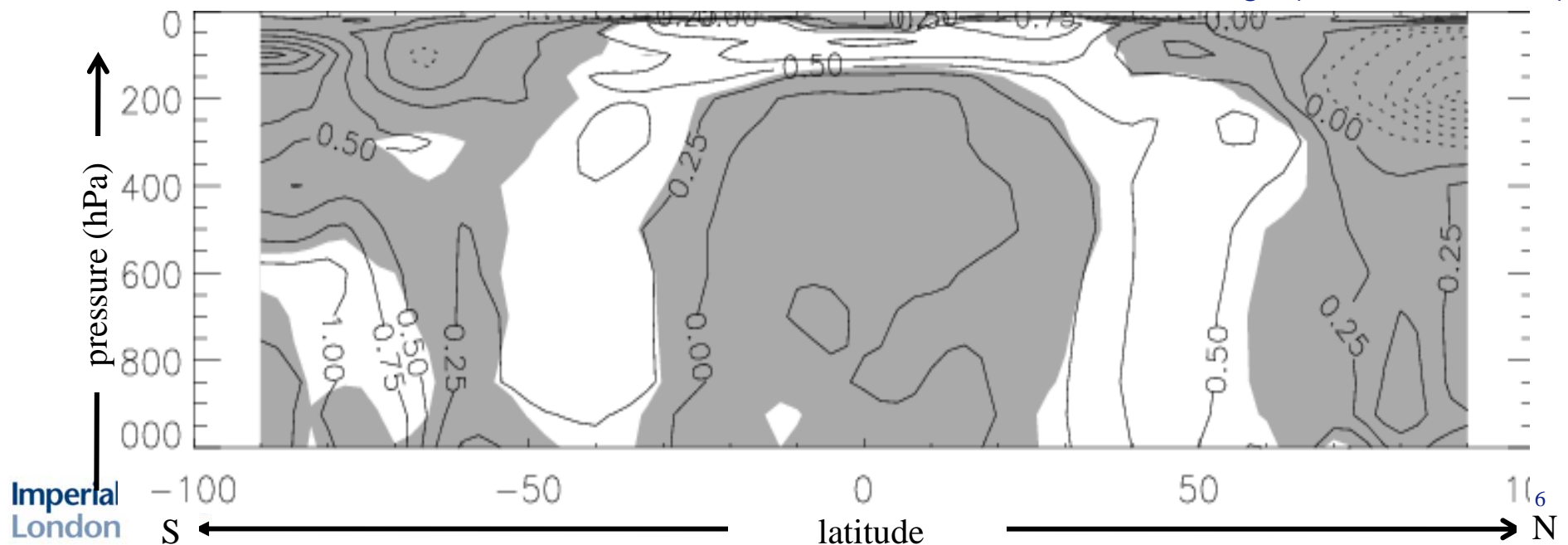
- Impact of solar variability on wind and circulation of troposphere: observations and models.
- Mechanism proposed via UV heating of stratosphere and dynamical coupling.
- Unusual behaviour of solar spectrum over declining phase of solar cycle 23.
- Implications for stratospheric composition and solar forcing of climate.

# Temperature changes over the 11-year solar cycle

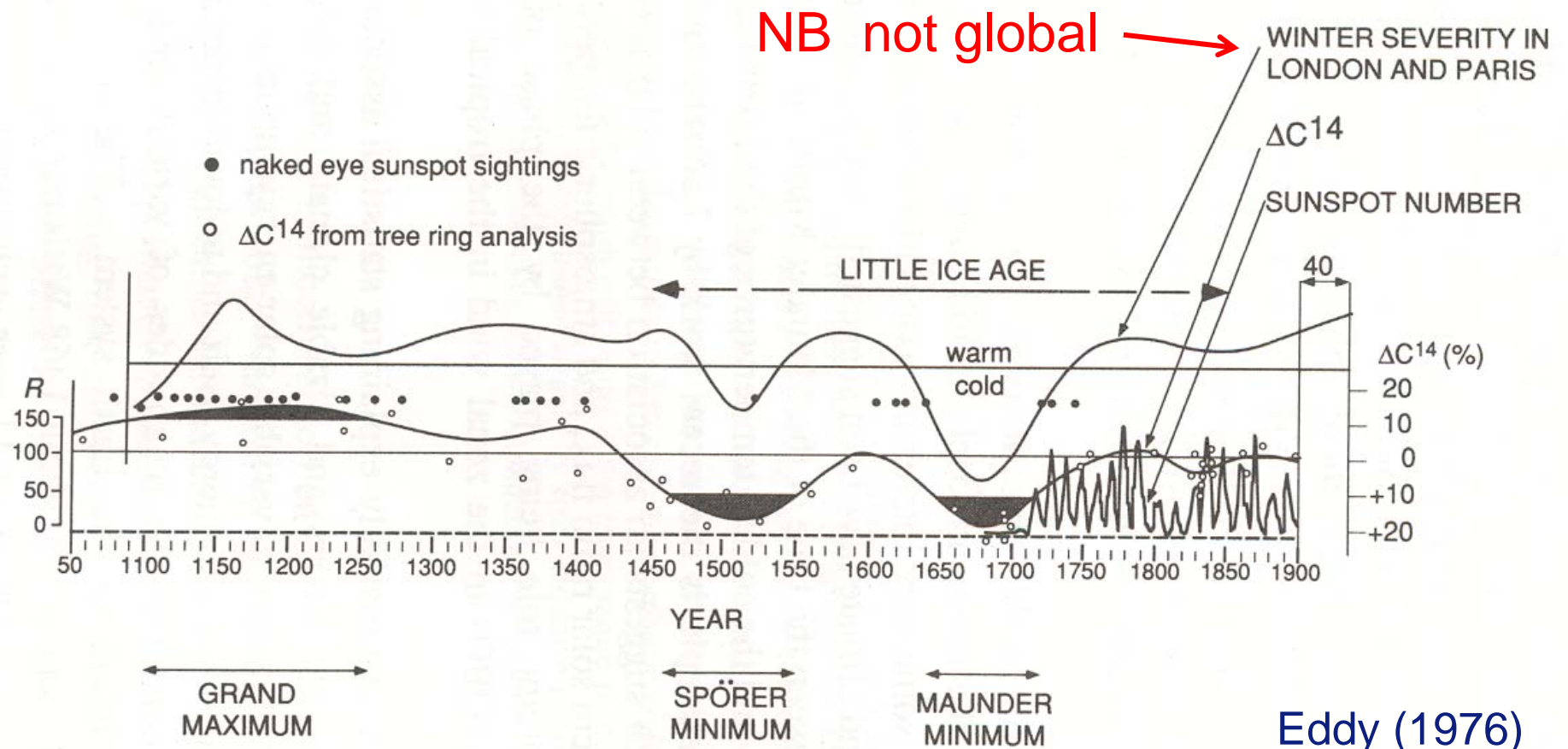
- Non-uniform.
- Warming in equatorial stratosphere, less towards the poles.
- Bands of warming in mid-latitudes.

Multiple regression analysis of zonal mean temperature 1979-2000

Haigh (Phil Trans, 2003)



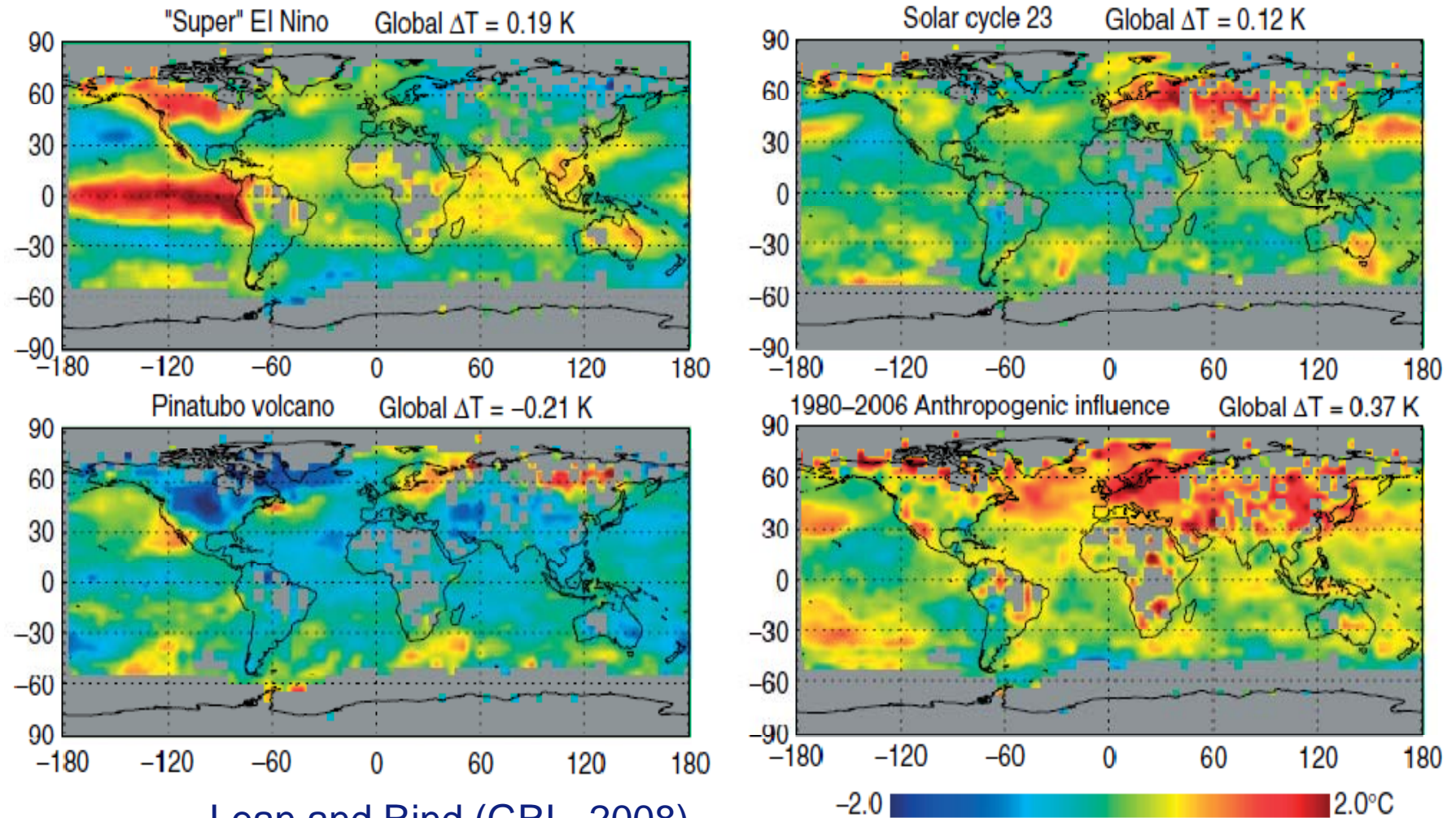
# Temperature in NW Europe



NB  $C^{14}$  produced by cosmic rays: more when Sun less active



# Signals in surface air temperature



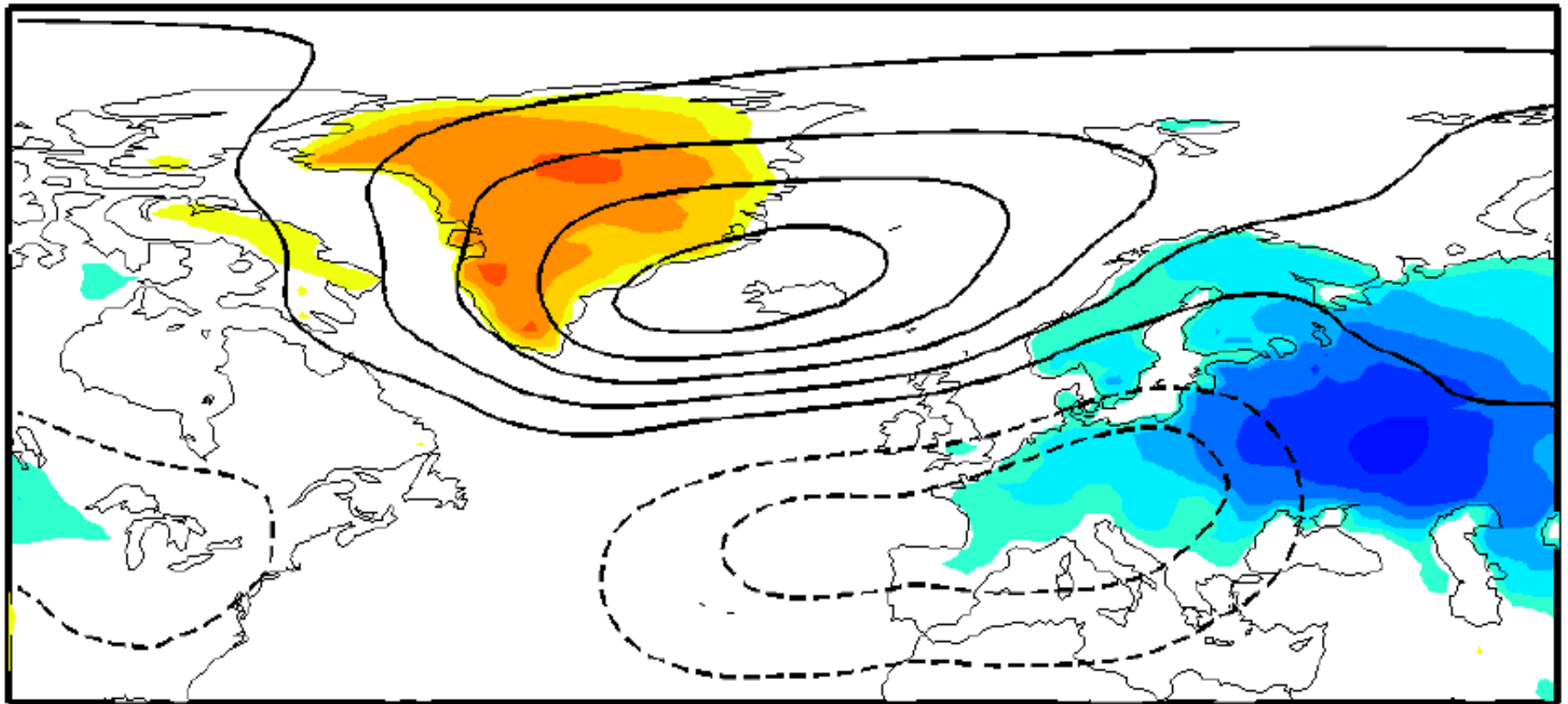
Lean and Rind (GRL, 2008)



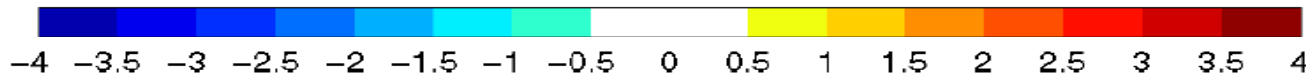
# Winter land surface air temperature and pressure

Lockwood et al. (ERL, 2010)

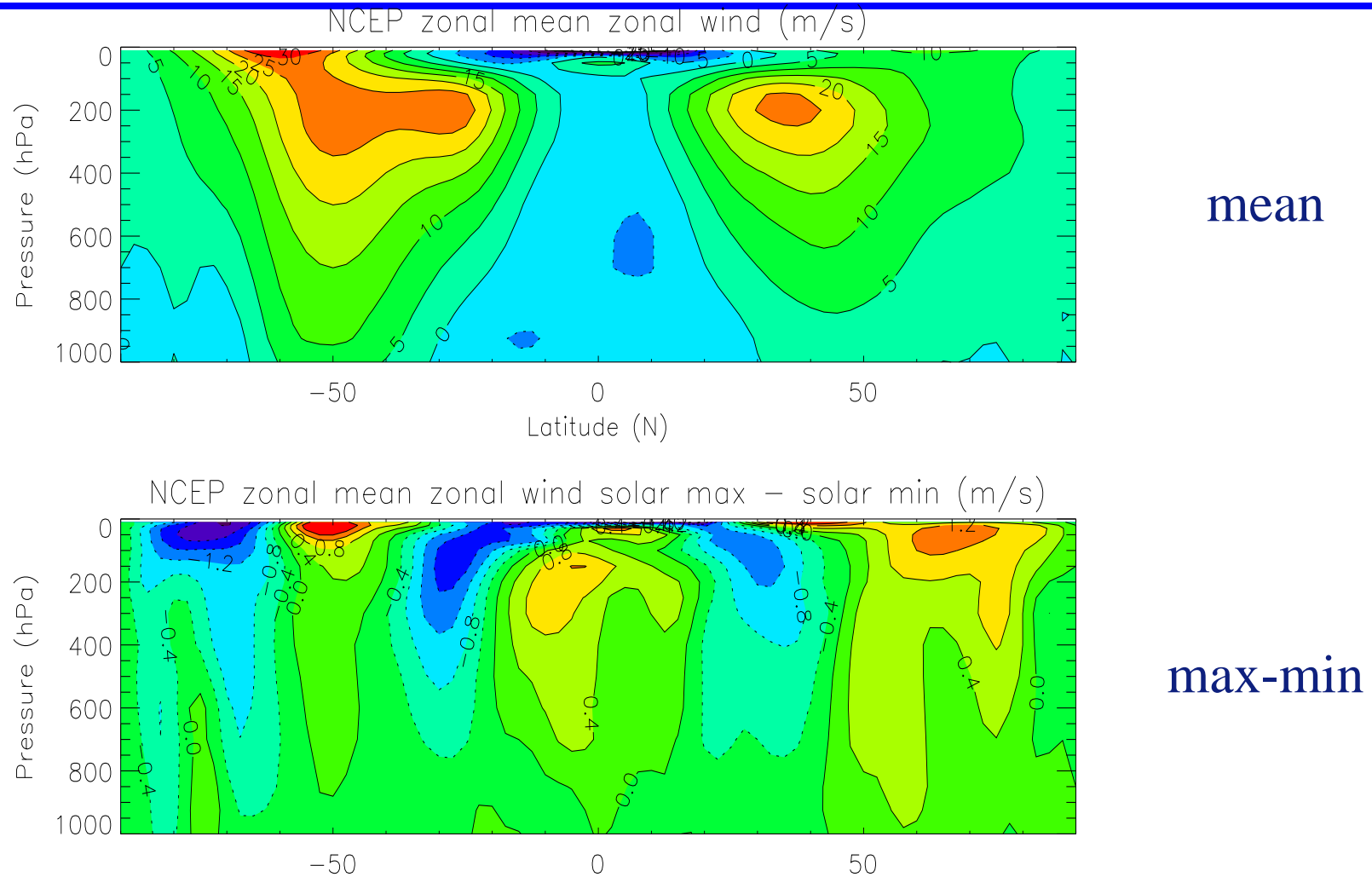
**SOLAR: LOW – HIGH**



2m Temperature (K)



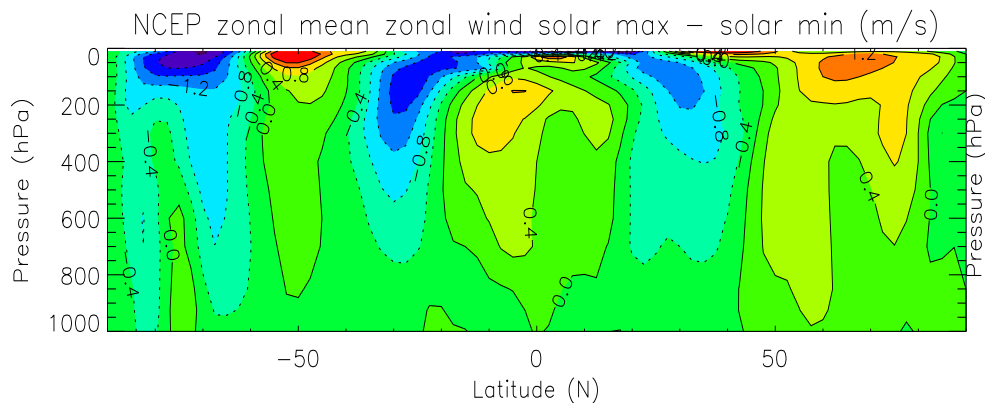
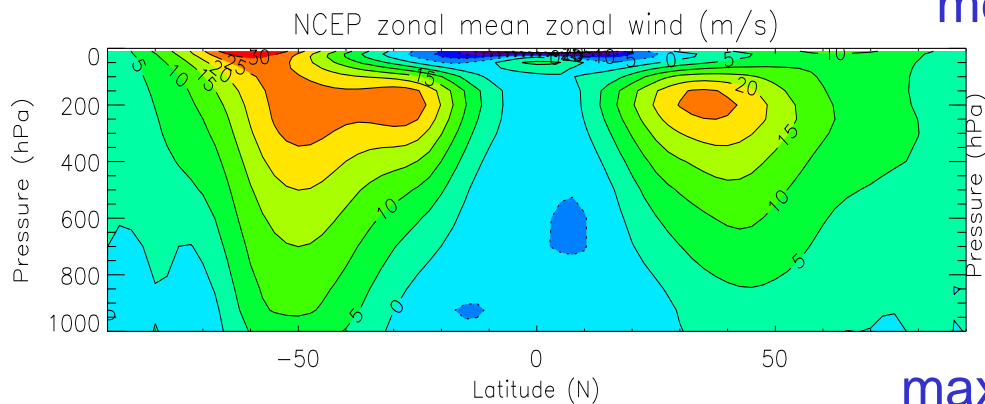
# Solar cycle signal in westerly wind (observed)



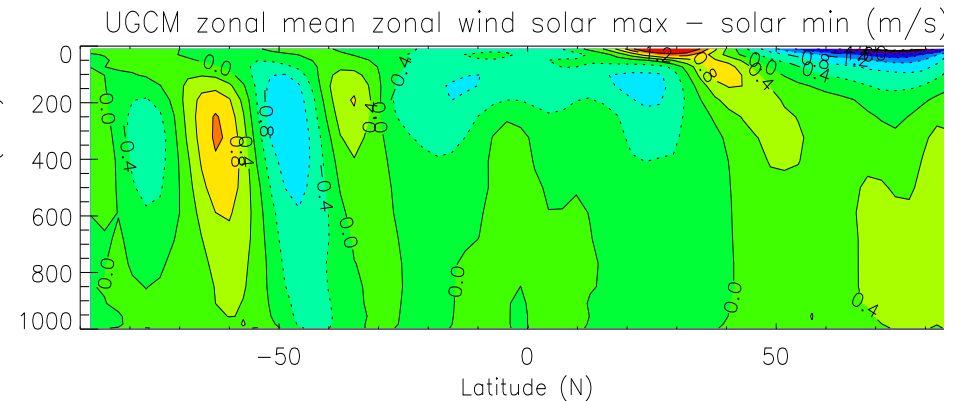
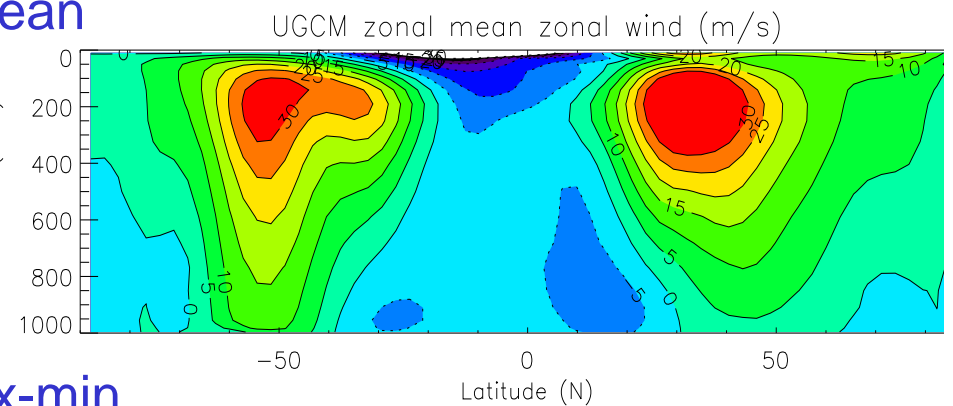
Weakening and poleward shift of the mid-latitude jets

# Solar signal in zonal wind: obs *cf* climate model

## Observations



## GCM response to changes in solar radiation & O<sub>3</sub>

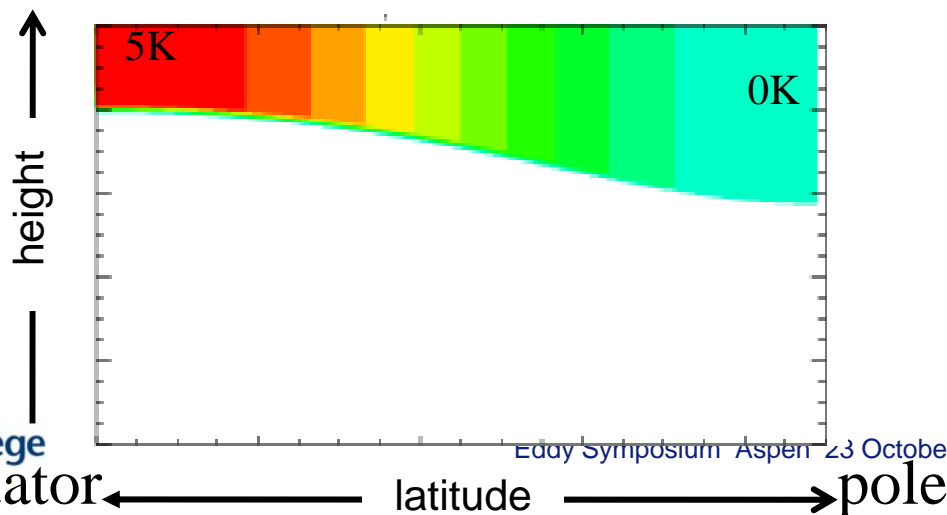


Haigh et al (J. Clim., 2005)

N.B. Similar pattern in response but much smaller amplitude without O<sub>3</sub> changes

# Study using a simplified climate model

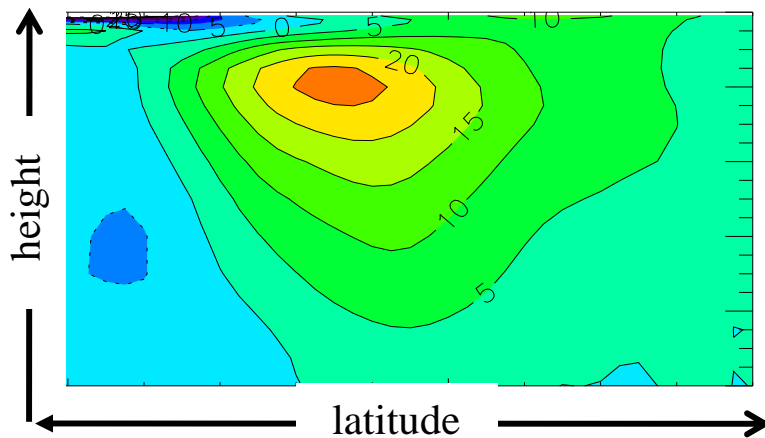
- Uses “Intermediate GCM”
  - full dynamics but highly simplified physics  
(so can try out ideas and do many runs cheaply)
  - no orography  
(so no planetary scale waves but still synoptic scale waves)
- Apply simplistic heating perturbation to the stratosphere:



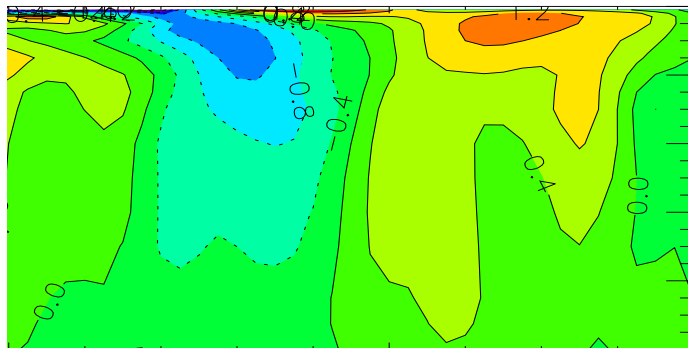
Haigh, Blackburn & Day (2005)

# Heating applied ONLY in the stratosphere:

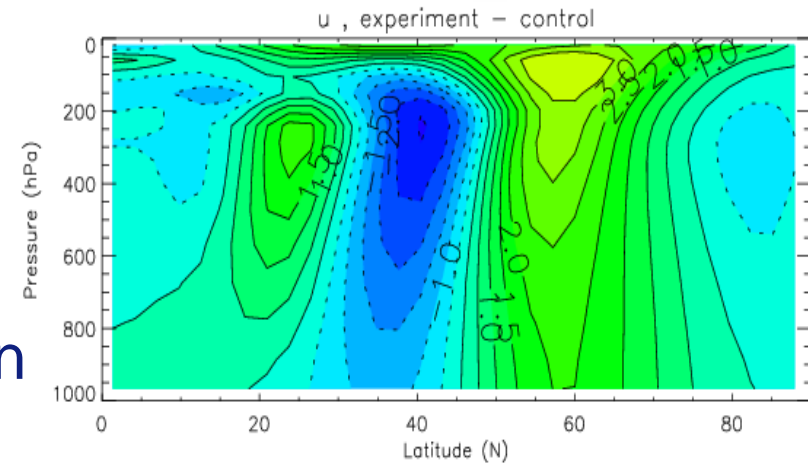
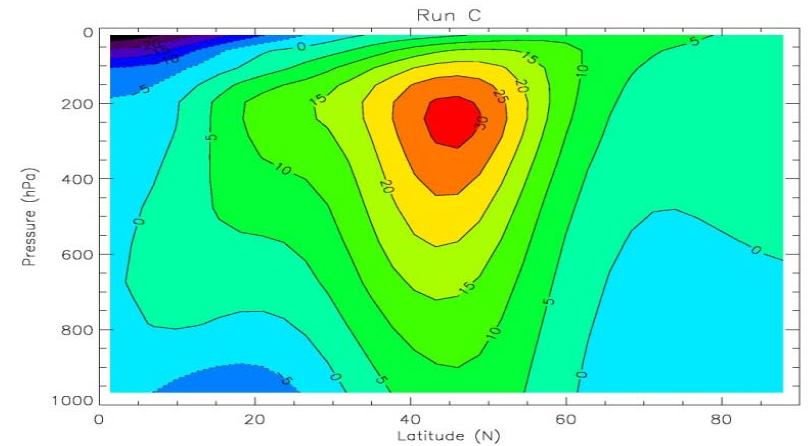
Observations of solar impact    zonal wind    Simple model



mean

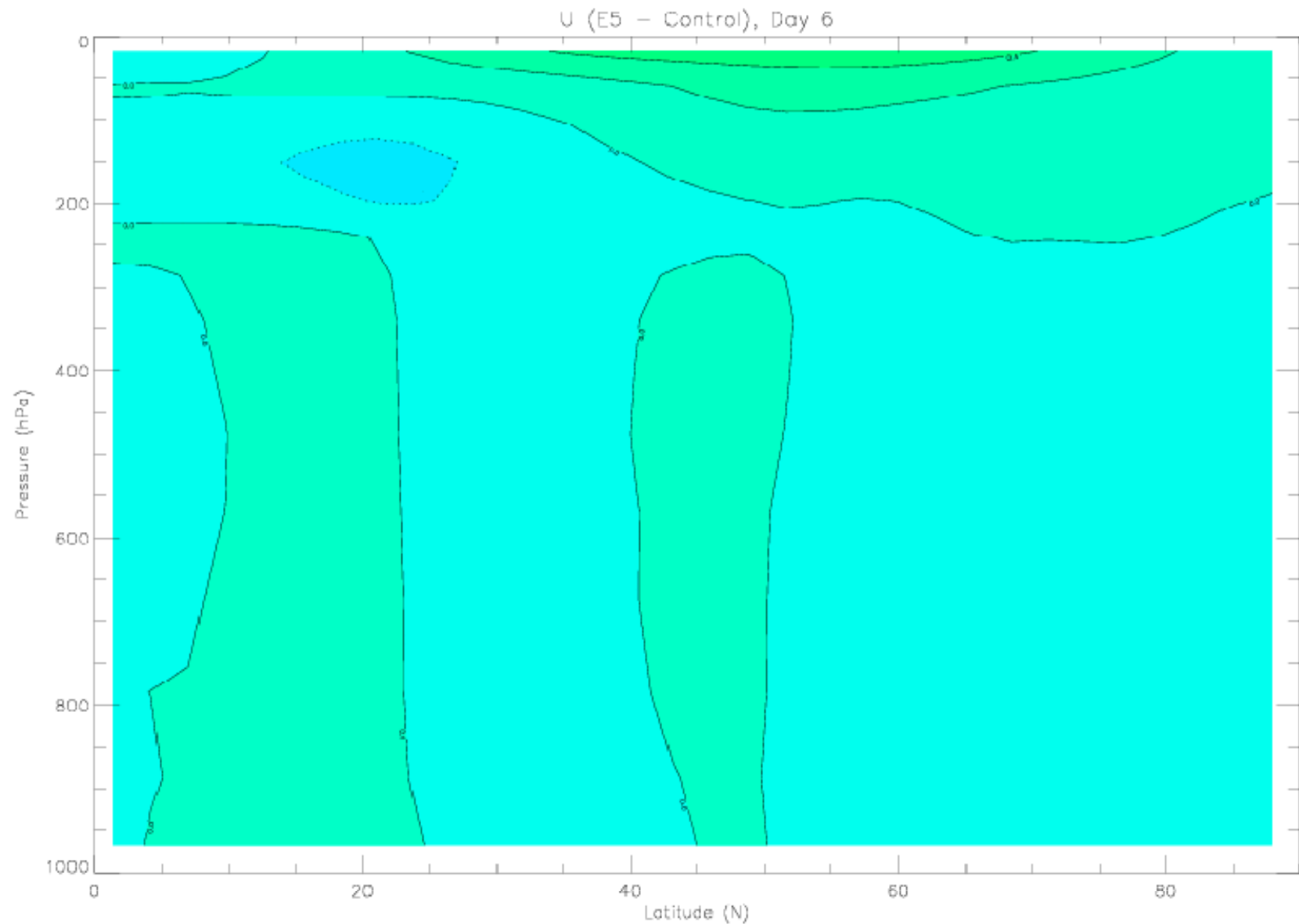


solar  
max - min



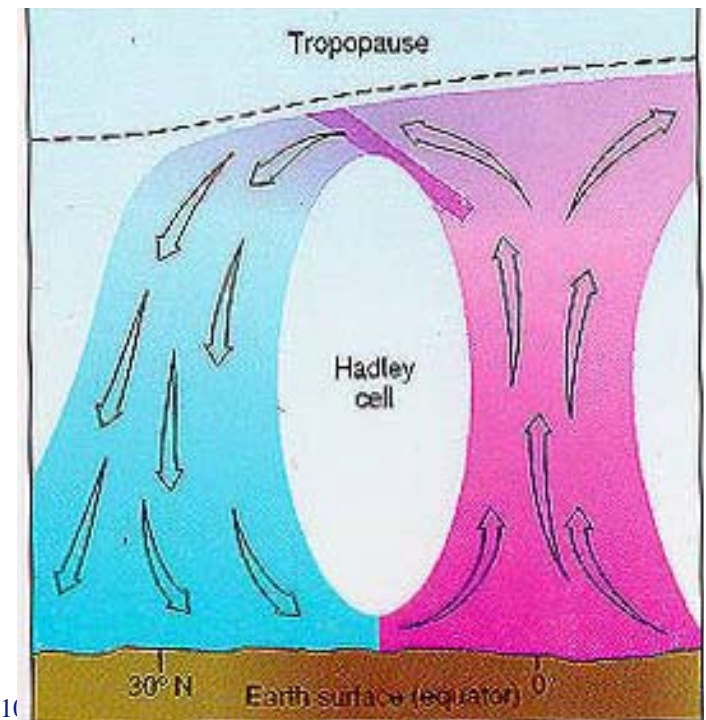
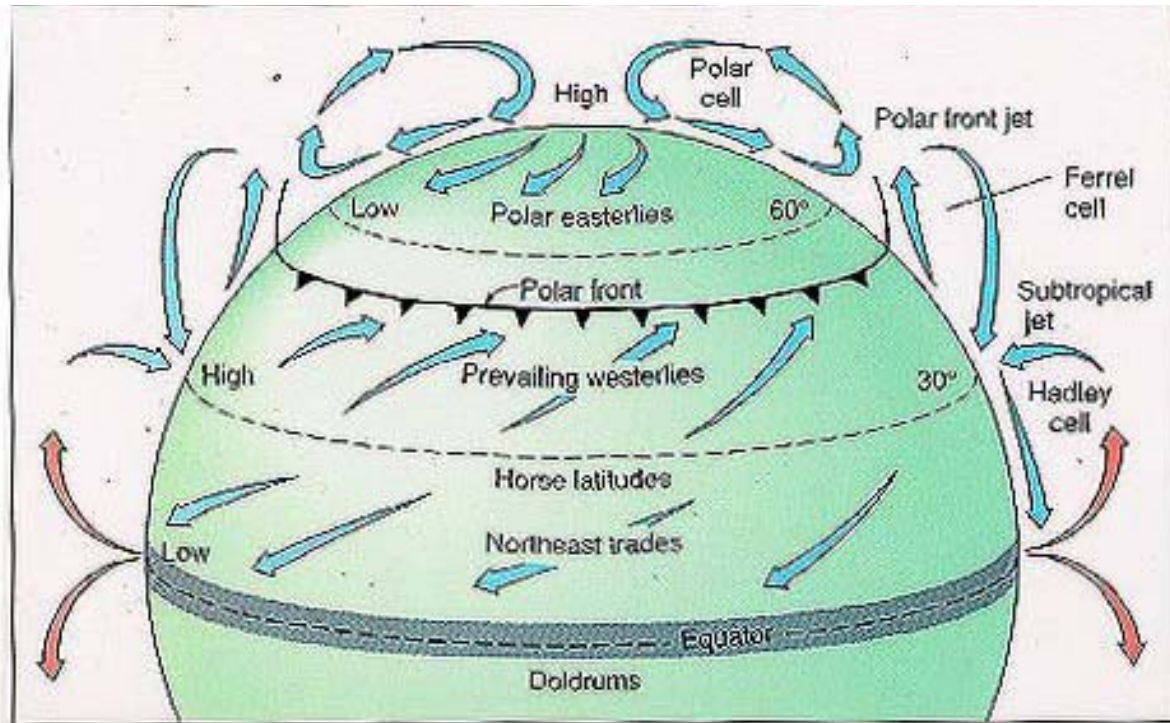
the jet weakens and moves polewards

# Evolution of zonal wind response 11-day running means





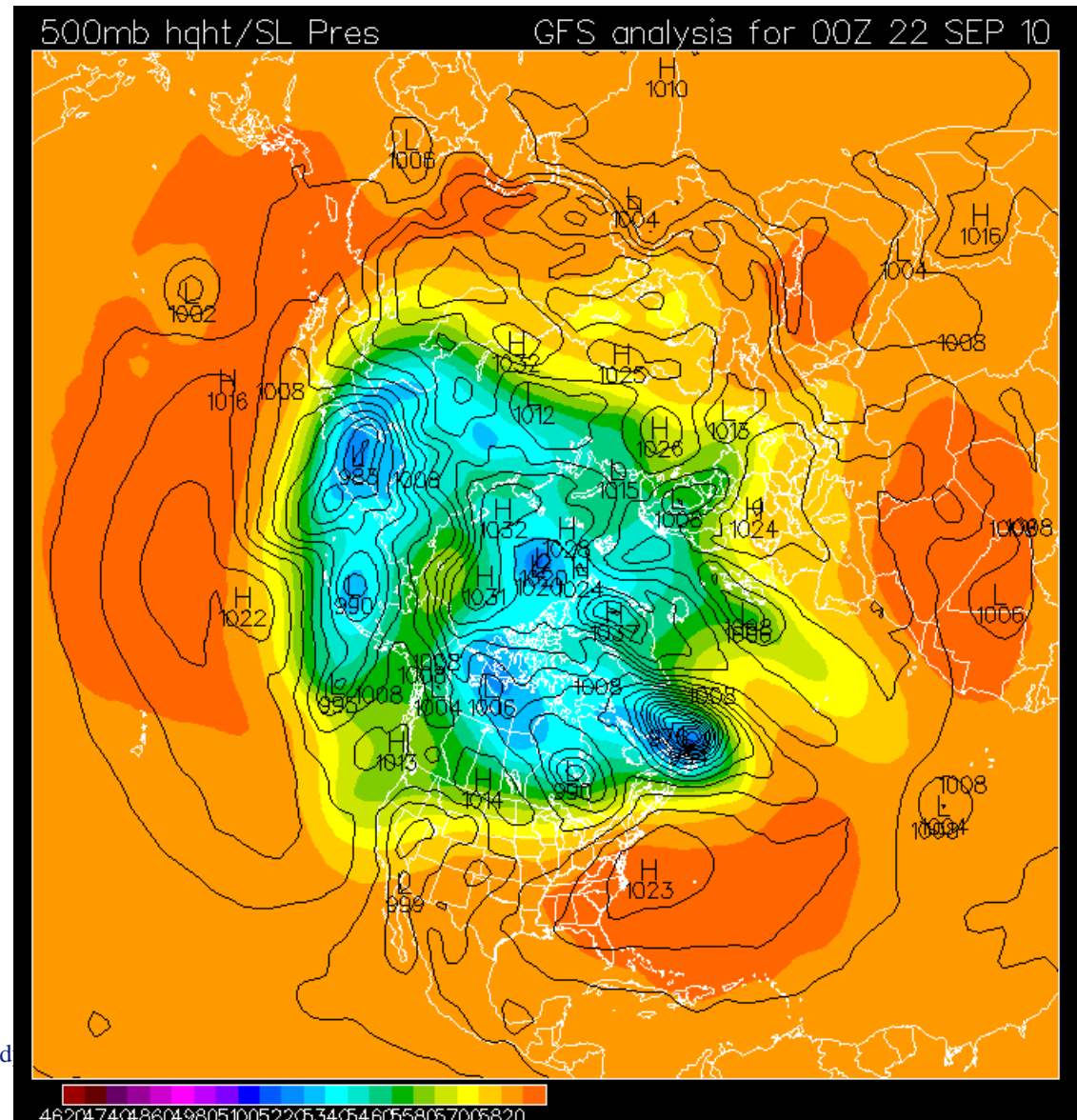
# Mean circulation of the tropical atmosphere: the Hadley cell



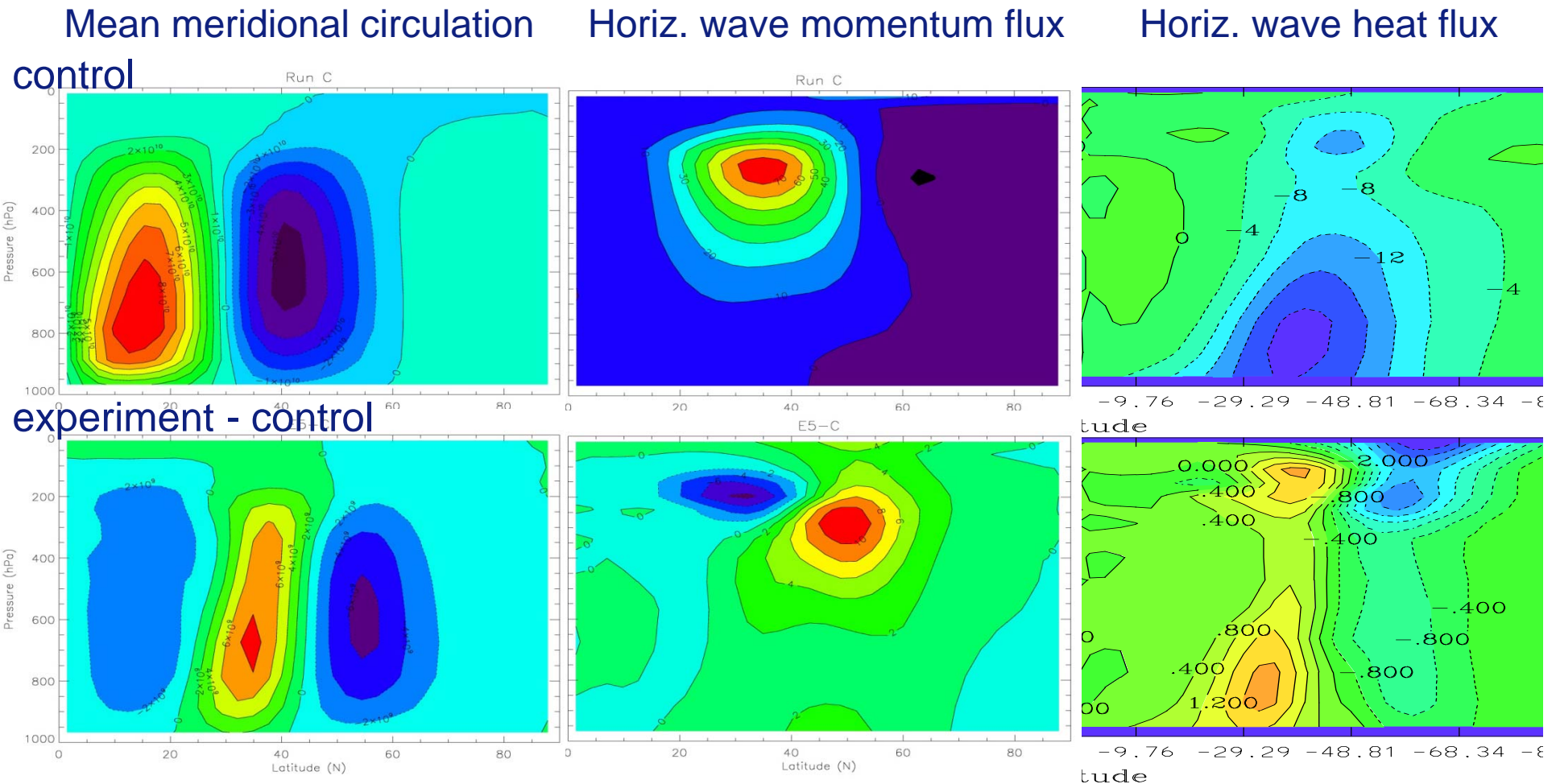
# Waves in the atmosphere

Contours: surface  
pressure (mb)

Colours: temperature  
of lowest 5.5km  
(approx)

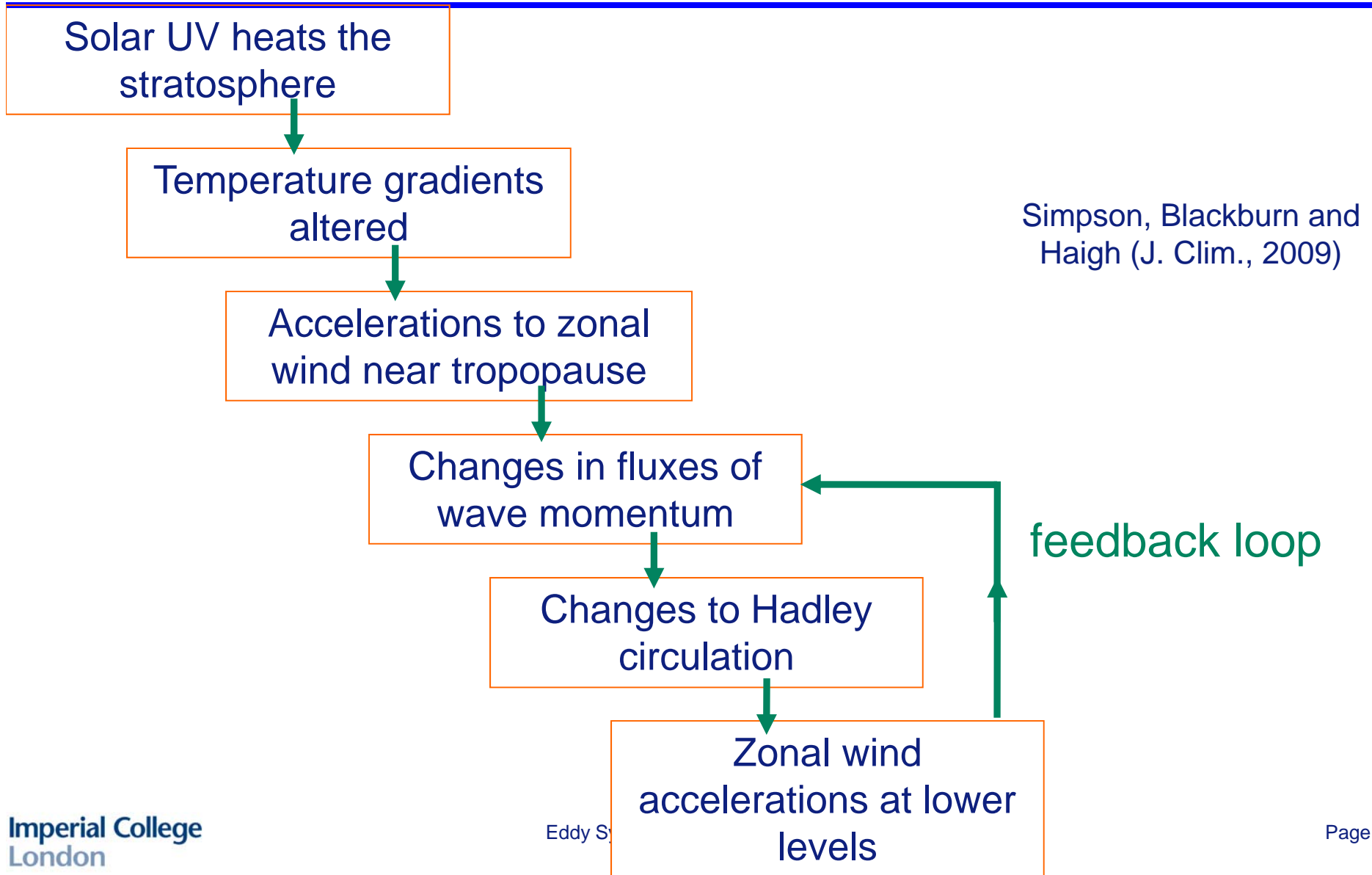


# Simple model equatorial heating (E5) results:



The Hadley cell weakens and expands; wave fluxes moves polewards

# Outline of mechanism:



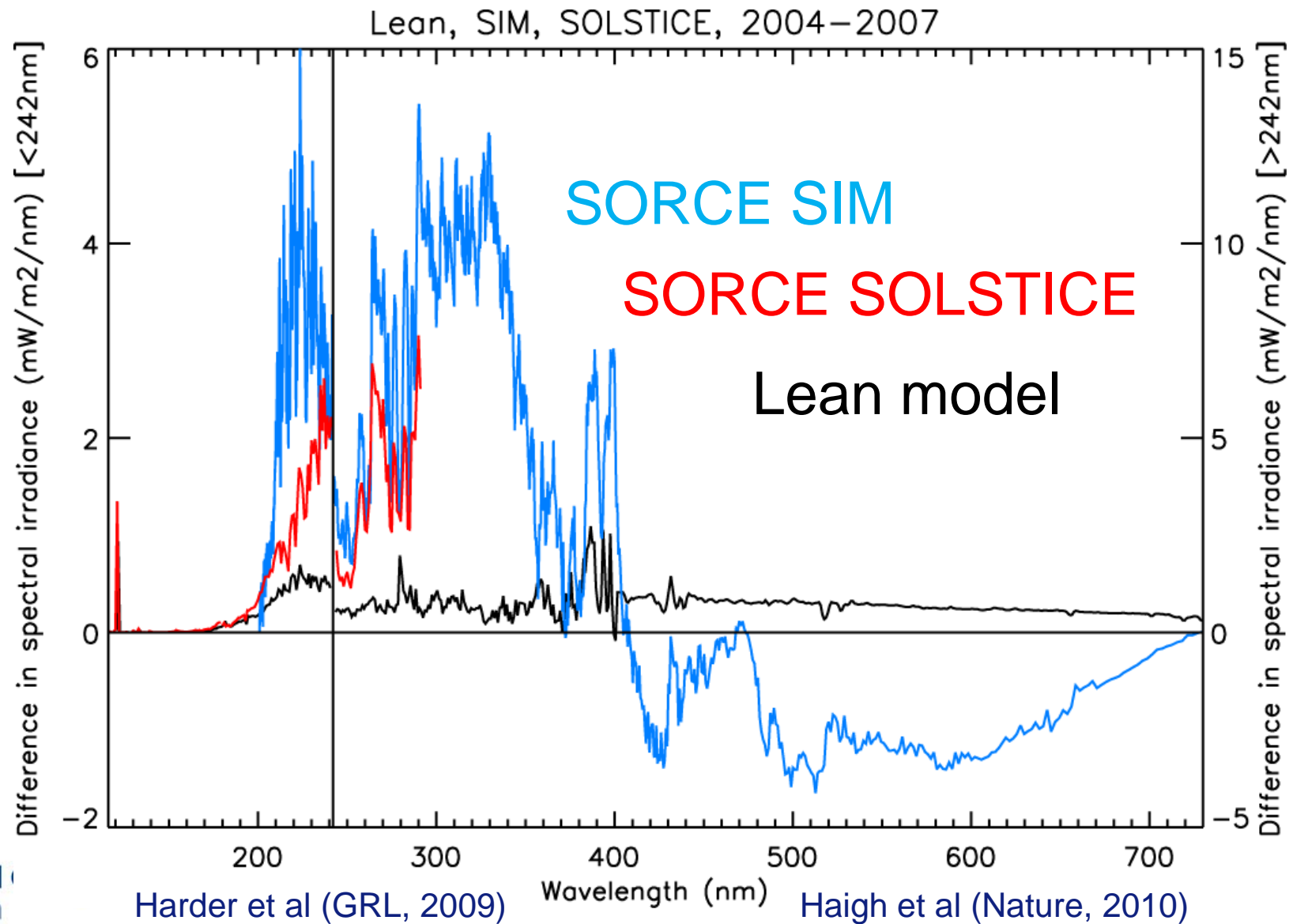
# Summary: Part I

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- (Solar) heating of the lower stratosphere produces changes in tropospheric circulation.
- Tropical heating produces a weakening and poleward shifts of the jets and a weakening and expansion of the Hadley cells.
- This results from the impact on the momentum budget of a feedback between vertically propagating, synoptic-scale waves and the mean flow.
- Crucially dependent on magnitude and location of stratospheric heating.



# Solar spectra: differences 2004-2007

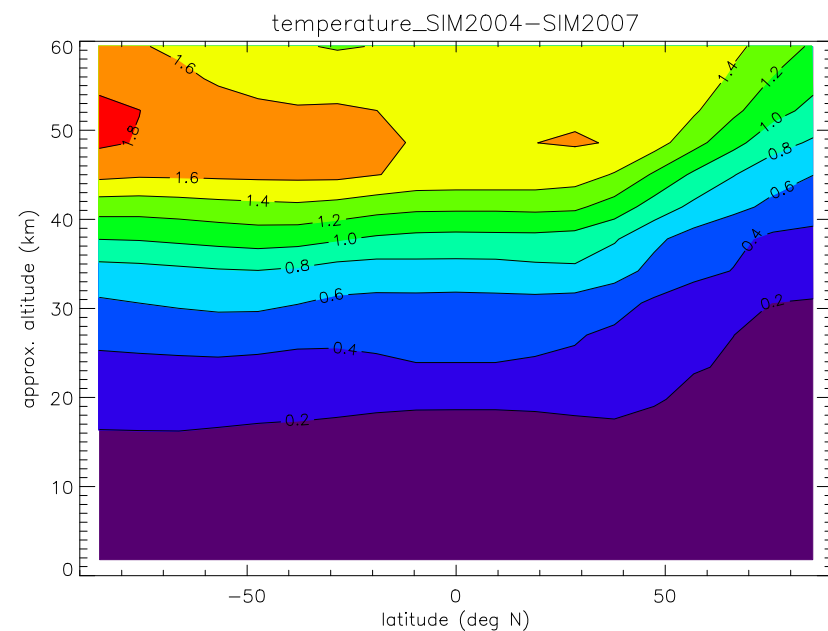
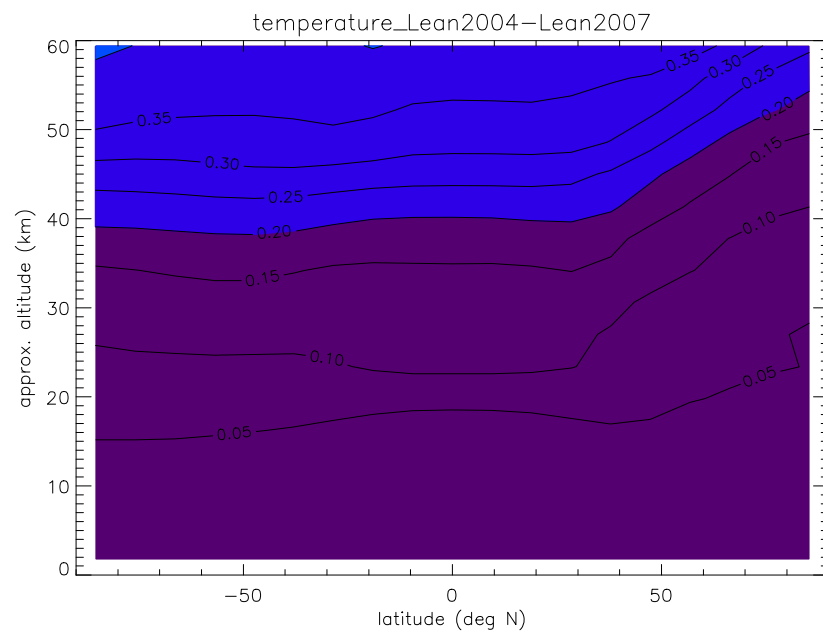




# 2D model temperature differences (K) 2004-2007

Lean

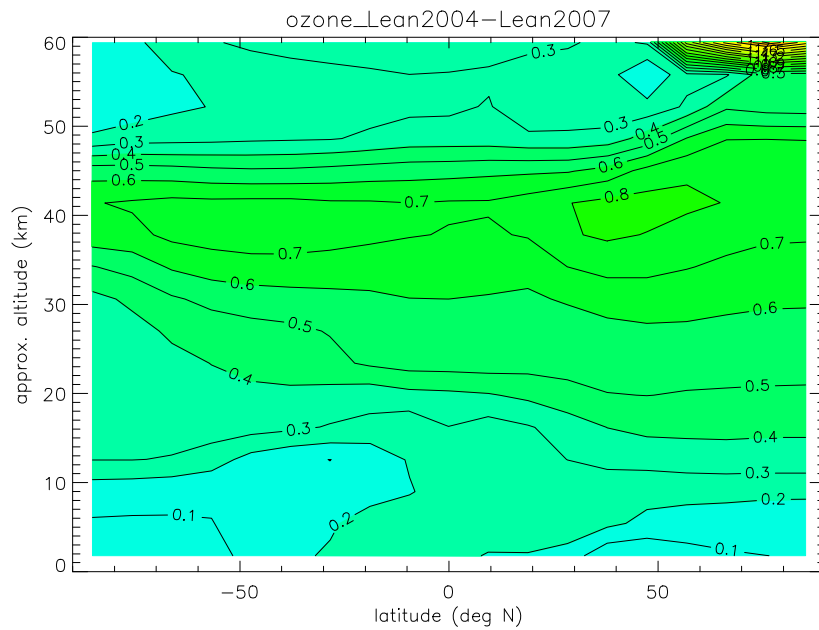
SIM



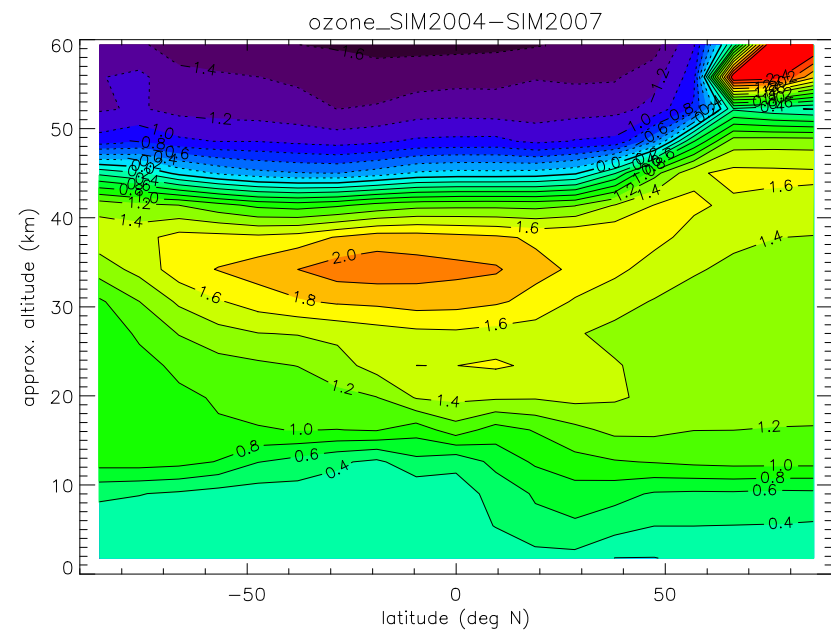
Haigh et al (Nature, 2010)

# 2D model O<sub>3</sub> differences (%) 2004-2007

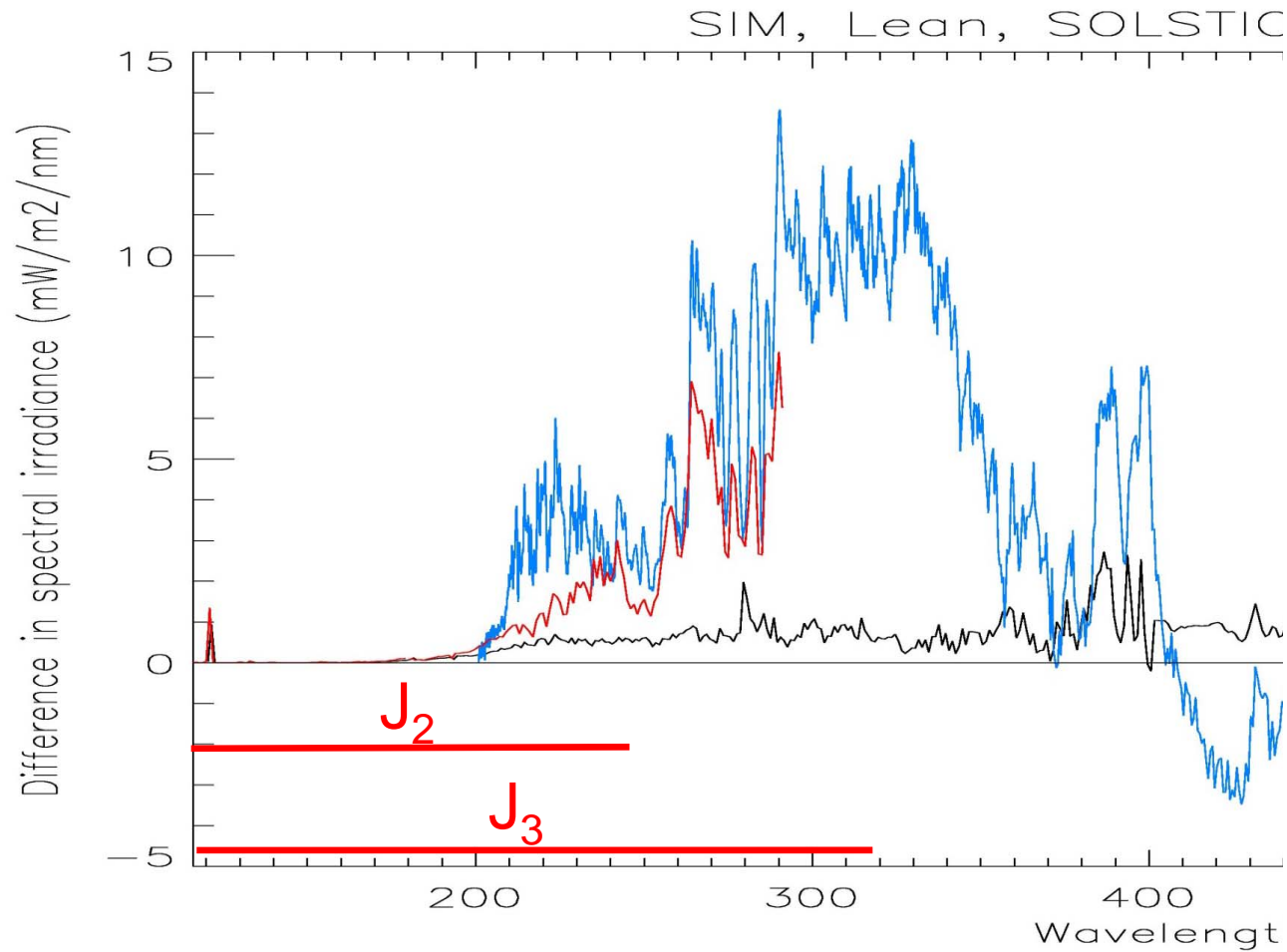
Lean



SIM



# Photodissociation



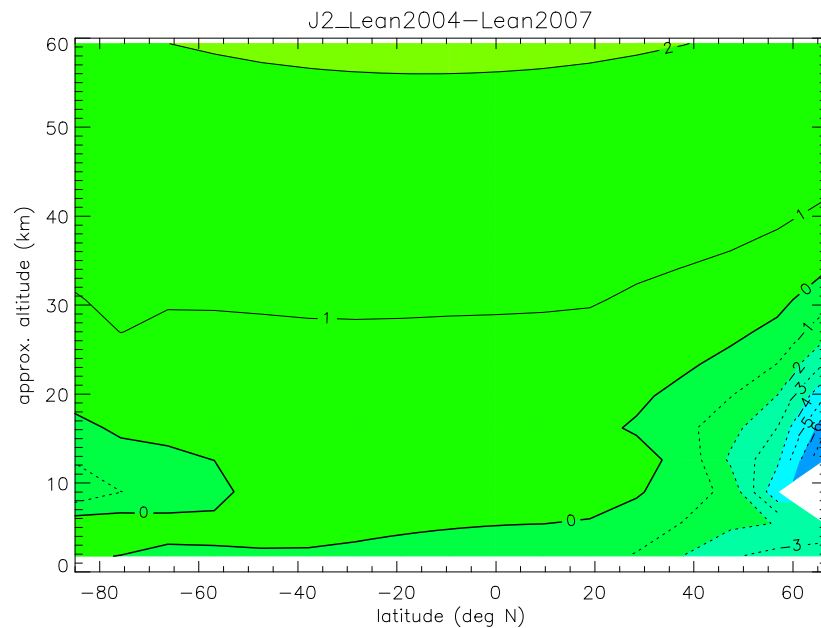
$J_2$  produces  
O and  $\text{O}_3$

$J_3$  produces  
 $\text{O}(^1\text{D})$  and  
determines  
 $\text{O}/\text{O}_3$  ratio

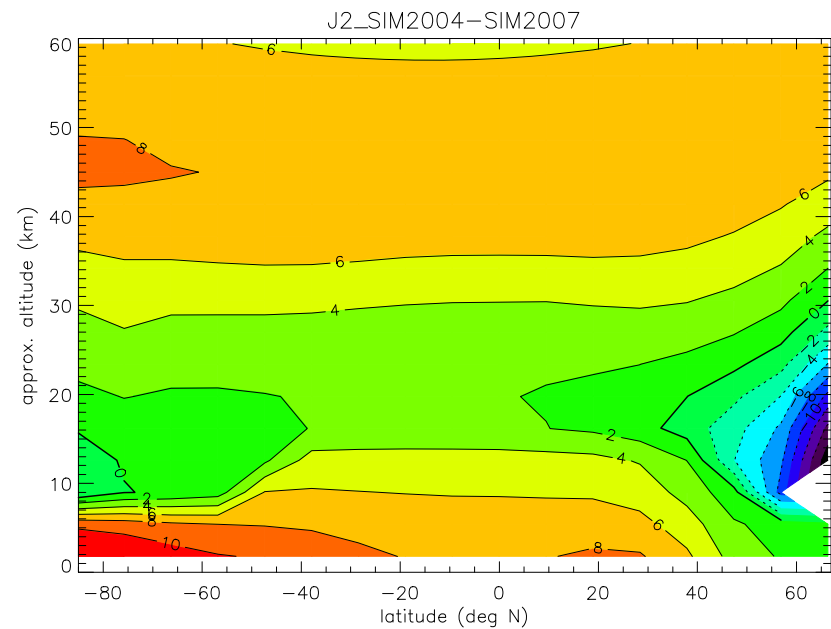
# $J_2$ : $O_2$ Photodissociation rate

2004 – 2007 (%)

Lean



SIM



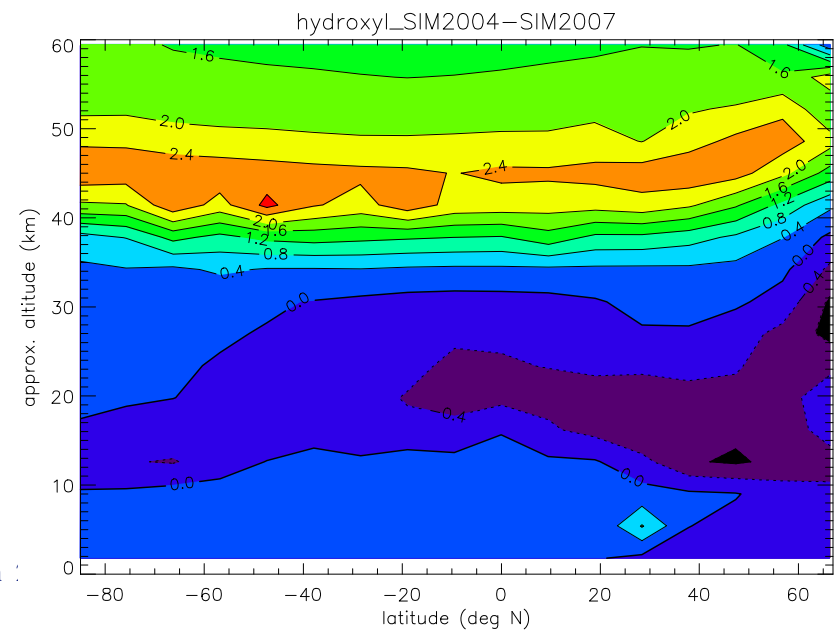
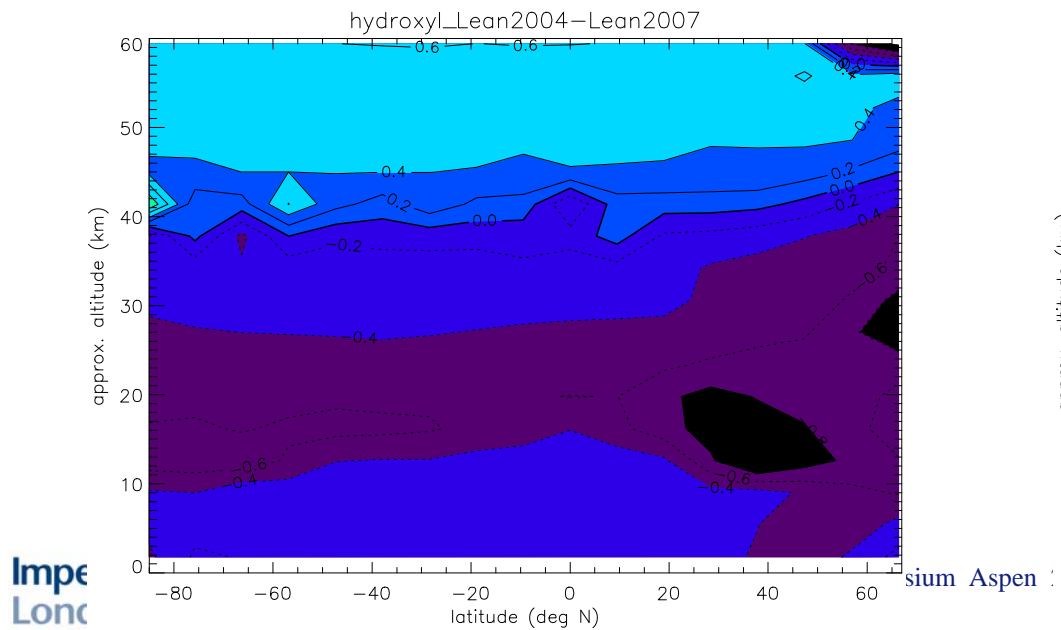
# OH

Produced by reaction of  $\text{H}_2\text{O}$  with  $\text{O}(^1\text{D})$

2004 – 2007 (%)

Lean

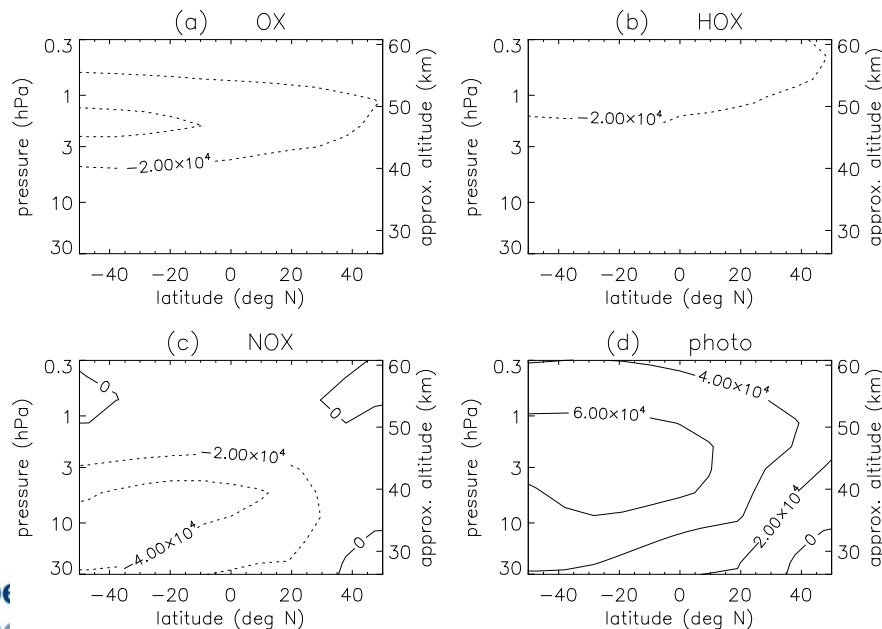
SIM



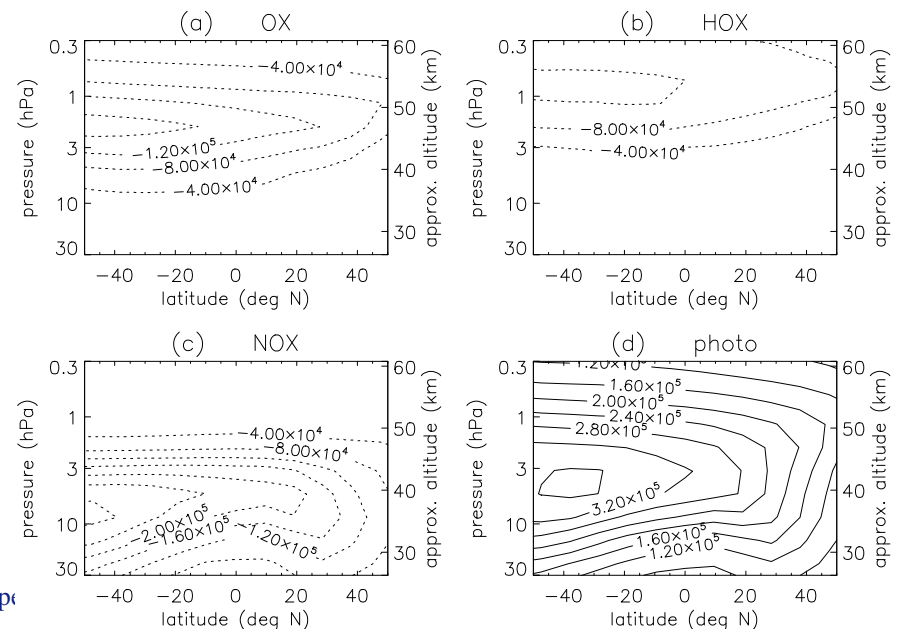
# O<sub>3</sub> production and destruction rates

- a) combination O+O<sub>3</sub>
- b) catalytic destruction by HO<sub>x</sub>
- c) catalytic destruction by NO<sub>x</sub>
- d) production by photodissociation of O<sub>2</sub>

Lean

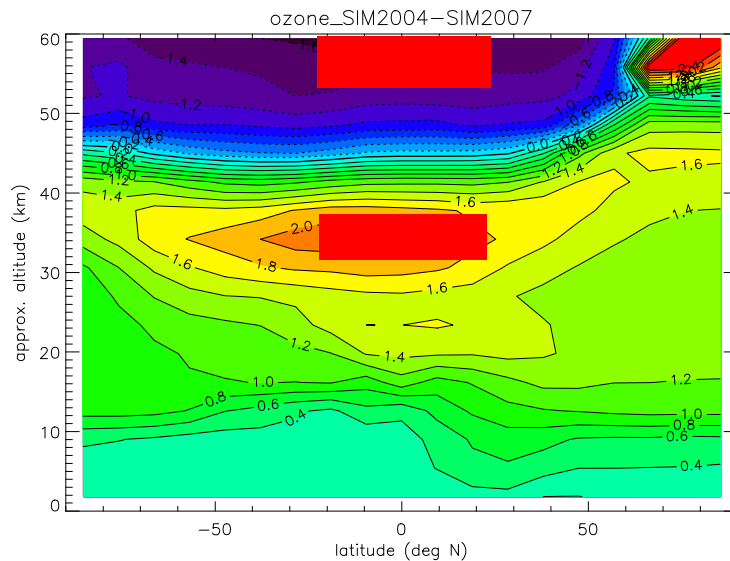


SIM





# Multiple regression of AURA MLS O<sub>3</sub> data



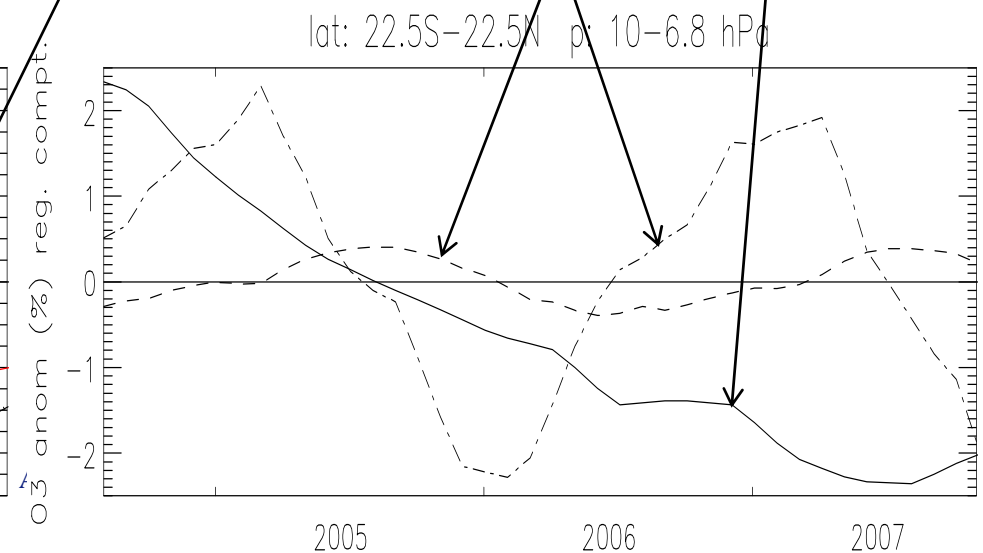
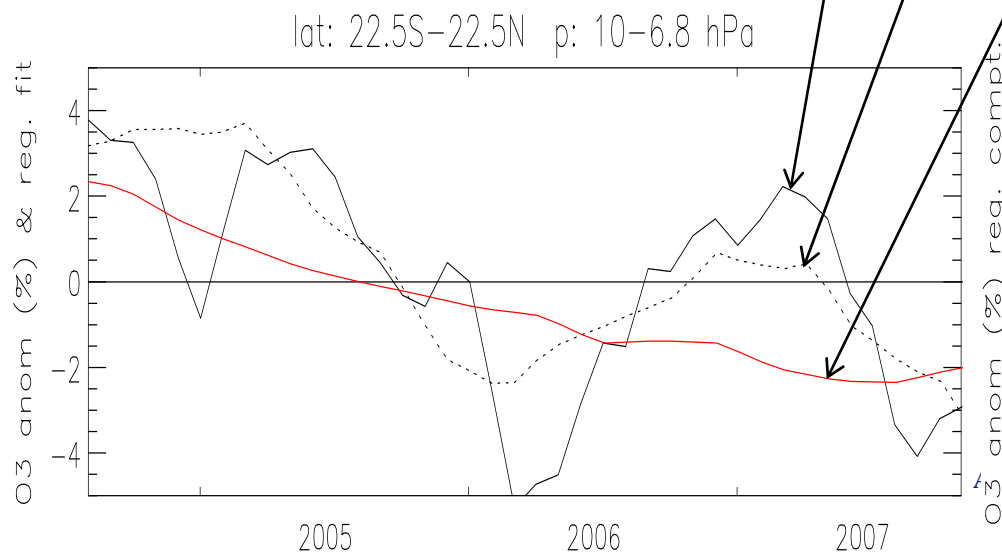
Regions chosen based on model results

Raw data (deseasonalised monthly means)  
for lower layer

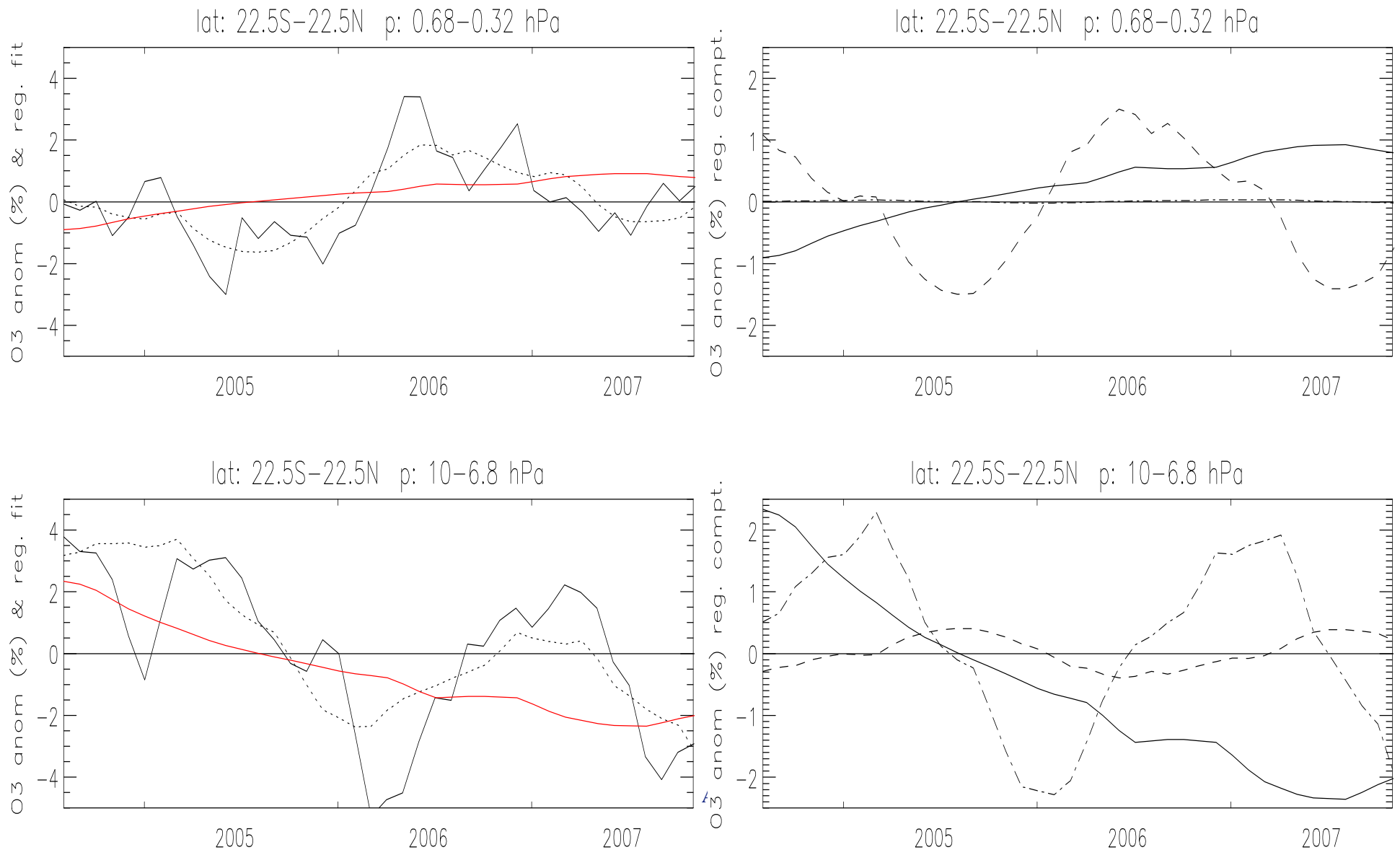
Regression fit

Regression indices:

Solar? constant, 2 QBO, SIM UV



# Multiple regression of AURA MLS O<sub>3</sub> data



# Integrated solar flux 2004-2007 ( $\text{Wm}^{-2}$ )

Lean

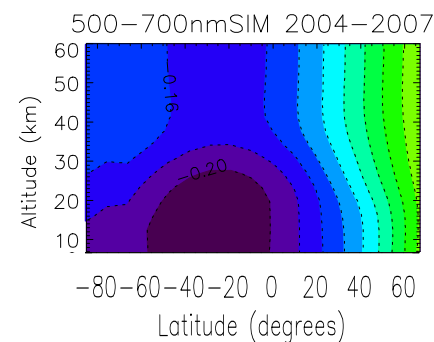
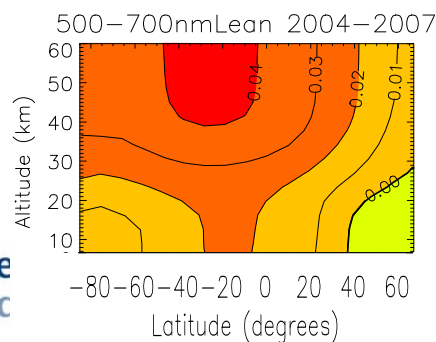
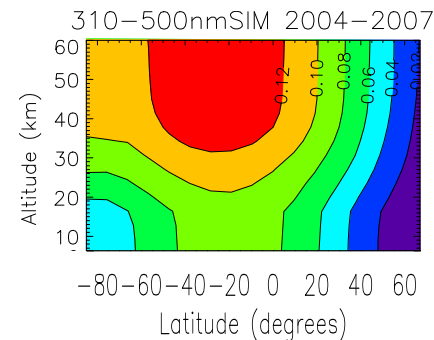
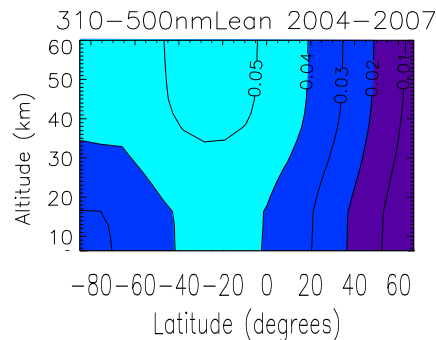
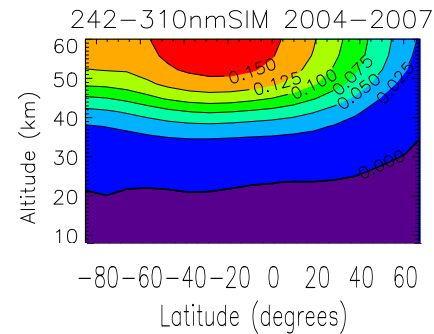
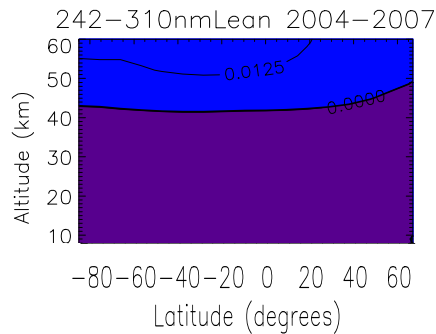
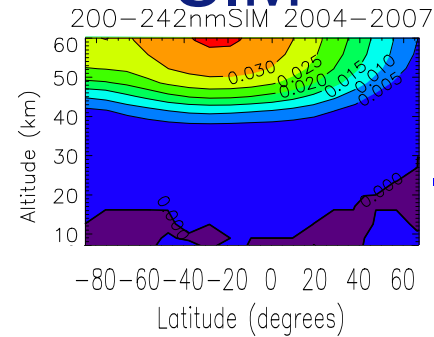
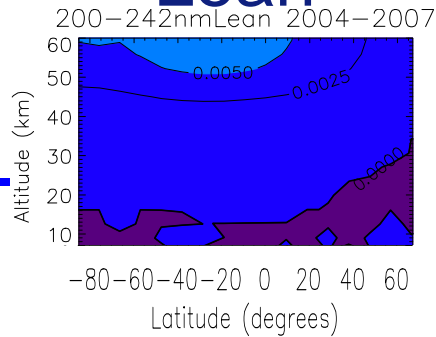
SIM

200-242 nm

242-310 nm

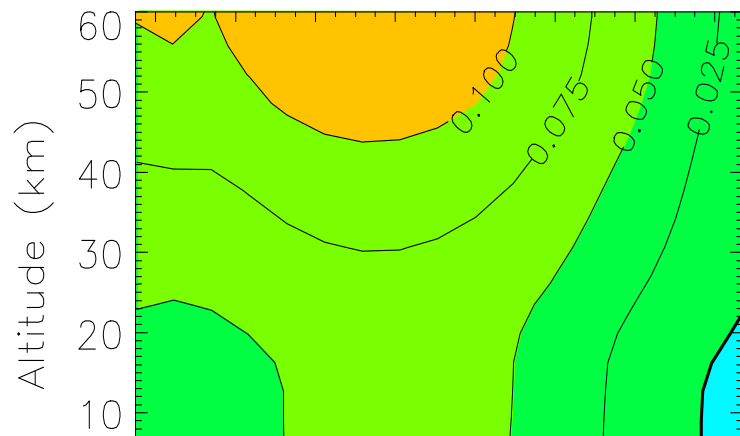
310-500 nm

500-700 nm

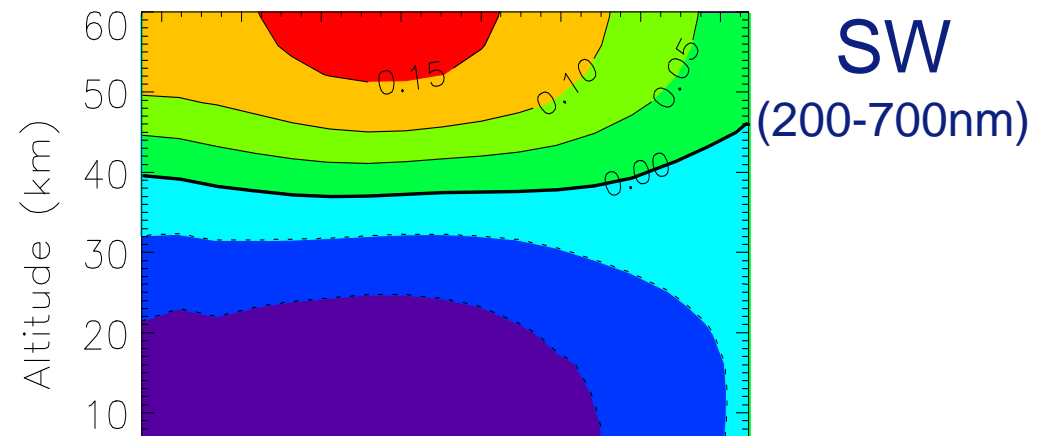


# Integrated radiative fluxes 2004-2007

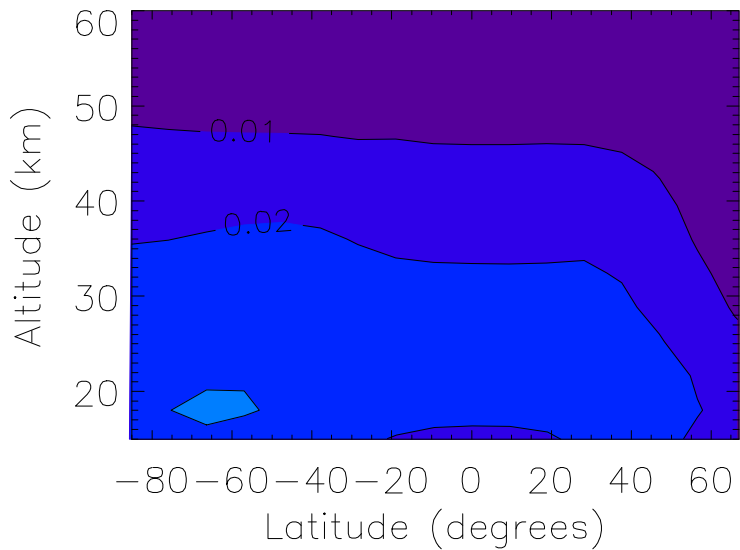
200–700nmLean 2004–2007



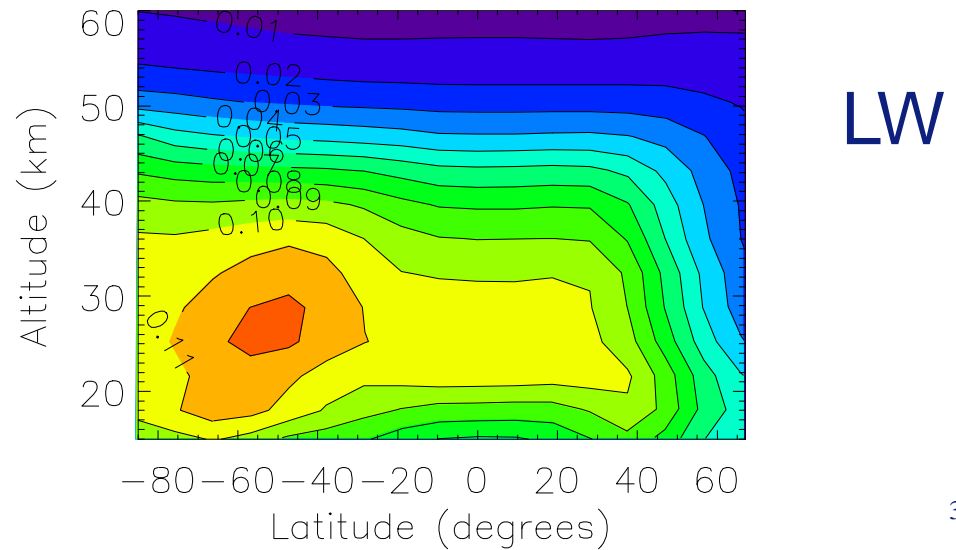
200–700nmSIM 2004–2007



downward LW Lean 2004–2007



downward LW SIM 2004–2007



# Solar Radiative Forcing of Climate\*

## 2004-2007 (mW m<sup>-2</sup>)

### Lean

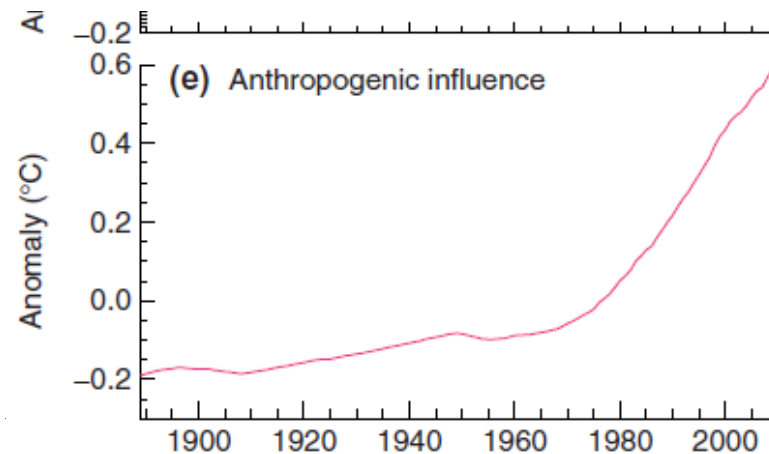
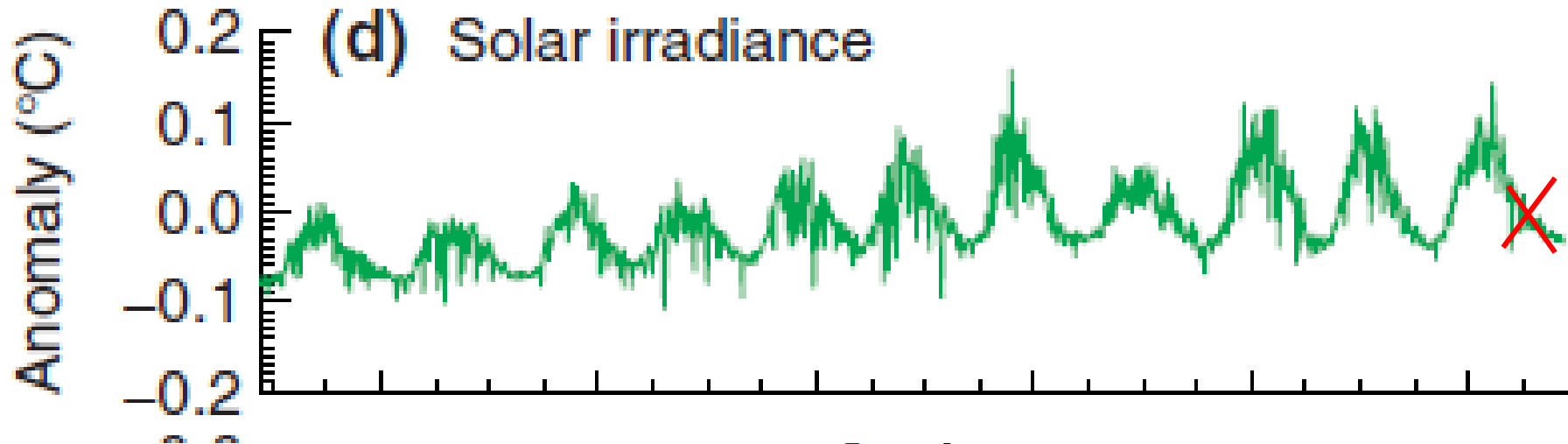
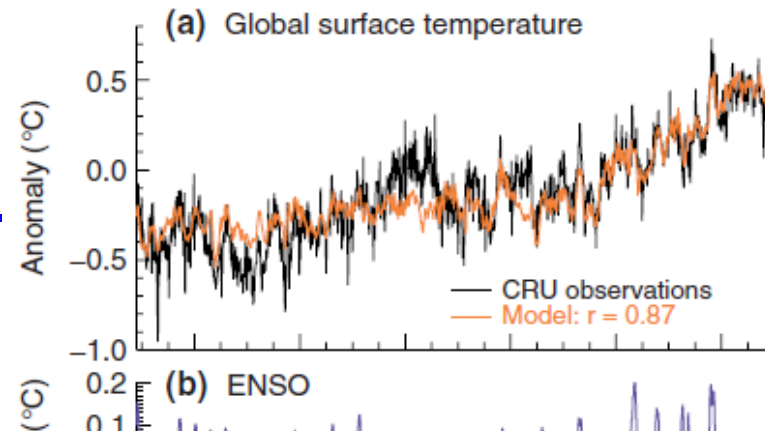
	200-310 nm	310-500 nm	500-700 nm	700-1600 nm	Total solar 200-1600 nm	Thermal (LW)	Net
TOA	20	40	30	20	110	0	110
105hPa	0	30	10	20	60	20	80

### SIM

	200-310 nm	310-500 nm	500-700 nm	700-1600 nm	Total solar 200-1600 nm	Thermal (LW)	Net
TOA	160	110	-130	-50	90	0	90
105hPa	0	60	-170	-50	-160	60	-100

# Global surface temperature

Influences of various



Lean  
(2010)



# Solar forcing of climate: current understanding

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“Top-down” via UV heating the stratosphere

New solar data would make much larger

and/or

“Bottom-up” via visible radiation warming surface ?

New solar data would invert

# Summary: Part 2

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- Input to 2D model SIM (& SOLSTICE) spectra produce a very different response in  $O_3$  from semi-empirical models of SSI: a *reduction* in lower mesosphere at higher solar activity and a large increase in mid- to upper stratosphere.
- This structure can be explained by enhanced production of  $HO_x$ , and by a shift of  $O_x$  from  $O_3$  to  $O$ .
- This structure is not inconsistent with contemporaneous measurements of  $O_3$  from AURA-MLS.
- SIM data would suggest that solar radiative forcing of climate produced a warming from 2004 to 2007, despite declining TSI.
- Is the Sun behaving oddly at present? Has this happened before?  
! What will the Sun do over the next few years...?

# SPARES

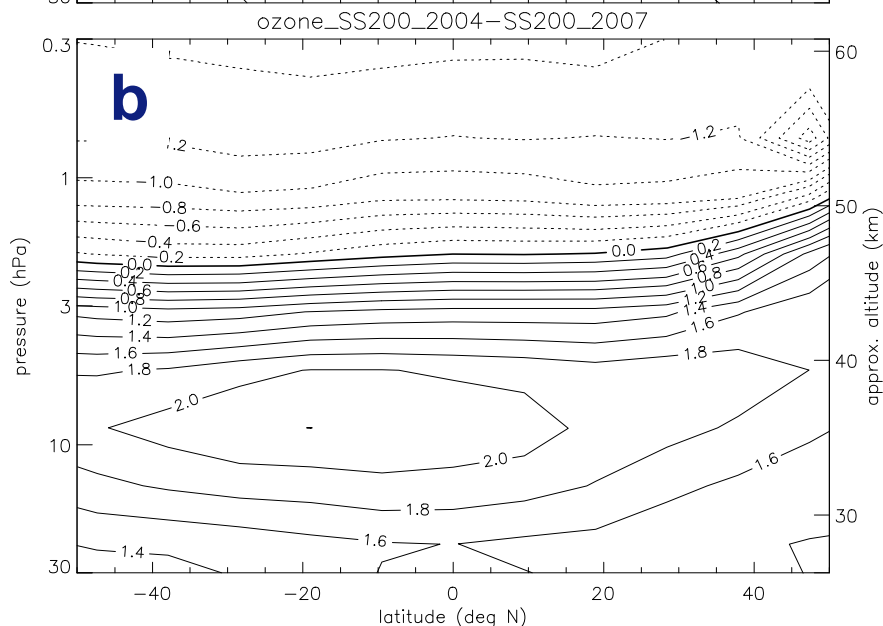
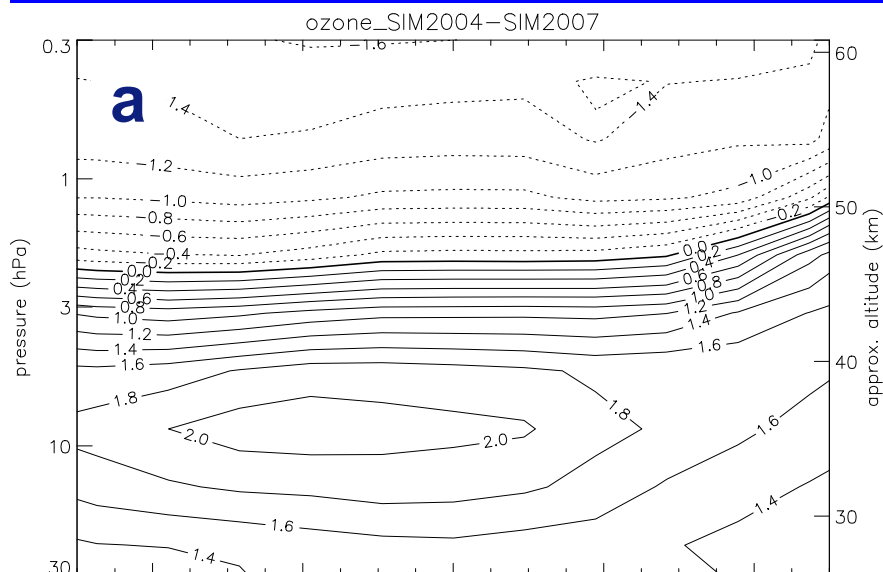
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# Choice of spectra $\lambda < 240$ nm

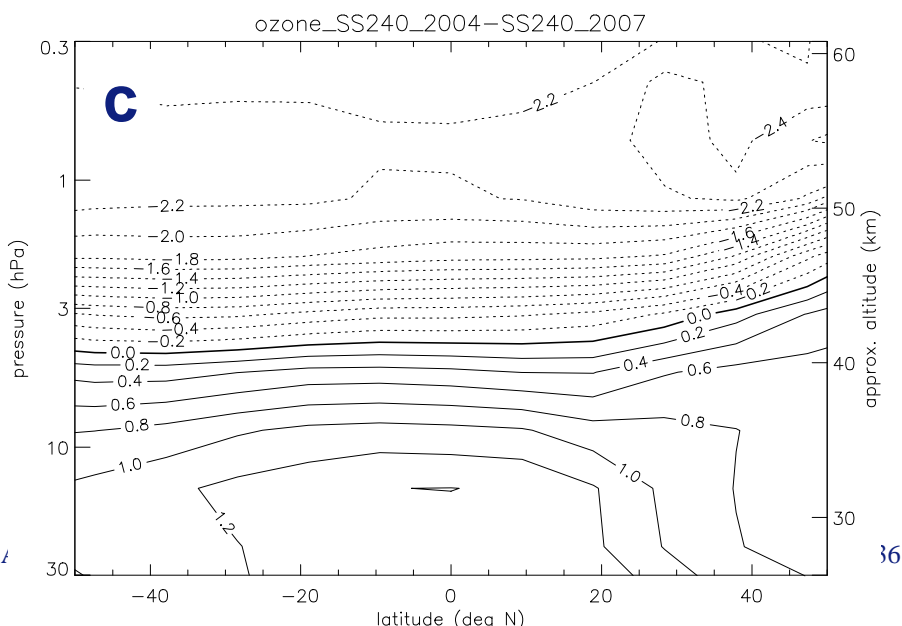
$\lambda < 200$      $200 < \lambda < 240$      $\lambda > 240$

<b>a</b> Lean	SIM	SIM
<b>b</b> SOLSTICE	SIM	SIM
<b>c</b> SOLSTICE	SOLSTICE	SIM

Ozone difference (%) 2004-2007



posium 4



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